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3.8 Changes in Extreme Events

3.8.1 Background

There is increasing concern that extreme events may be changing in frequency and intensity as a result of human influences on climate. Climate change may be perceived most through the impacts of extremes, although these are to a large degree dependent on the system under consideration, including its vulnerability, resiliency and capacity for adaptation and mitigation; see the Working Group II contribution to the IPCC Fourth Assessment Report. Improvements in technology mean that people hear about extremes in most parts of the world within a few hours of their occurrence. Pictures shot by

camcorders on the news may foster a belief that w extremes are increasing in frequency, whether the An extreme weather event becomes a disaster who or ecosystems are unable to cope with it effective human vulnerability (due to growing numbers of in exposed and marginal areas or due to the demore high-value property in high-risk zones) is risk, while human endeavours (such as by local try to mitigate possible effects.

IPCC – AR4 (2007)

extremes refer to rare events based on a statistical model of particular weather elements, and changes in extremes may relate to changes in the mean and variance in complicated ways. Changes in extremes are assessed at a range of temporal and spatial scales, for example, from extremely warm years globally to peak rainfall intensities locally, and examples are given in Box 3.6. To span this entire range, data are required at a daily (or shorter) time scale. However, the availability of observational data restricts the types of extremes that can be analysed. The rarer the event, the more difficult it is to identify long-term changes, simply because there are fewer cases to evaluate (Frei and Schär, 2001; Klein Tank and Können, 2003).

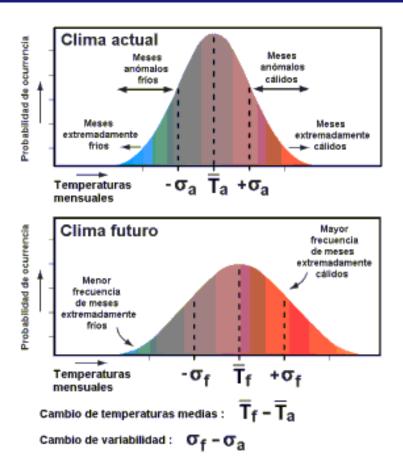


Figura 16. Ilustración esquemática de las distribuciones de frecuencia de las temperaturas mensuales en clima actual y en el clima proyectado.

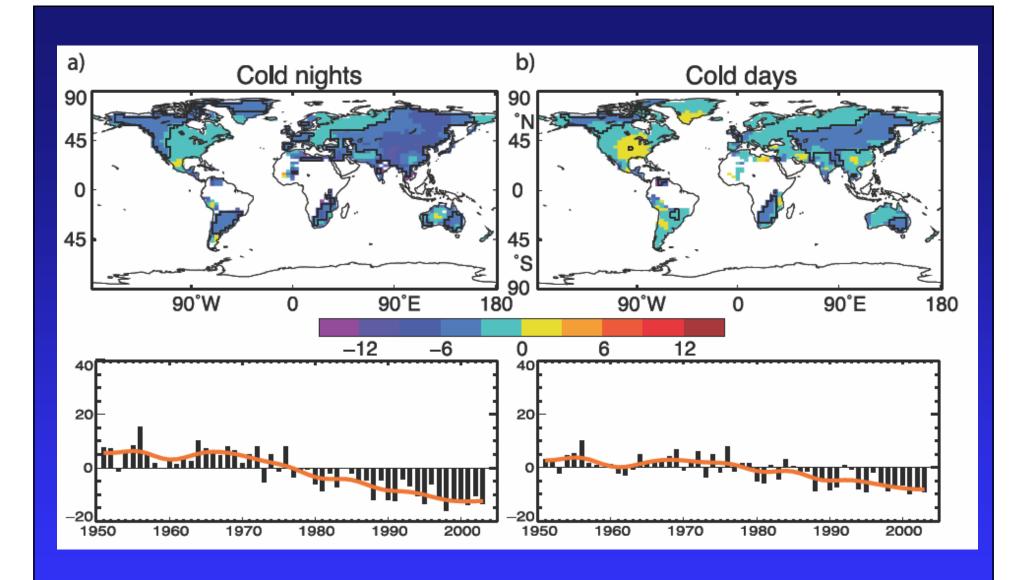
Frequently Asked Question 3.3

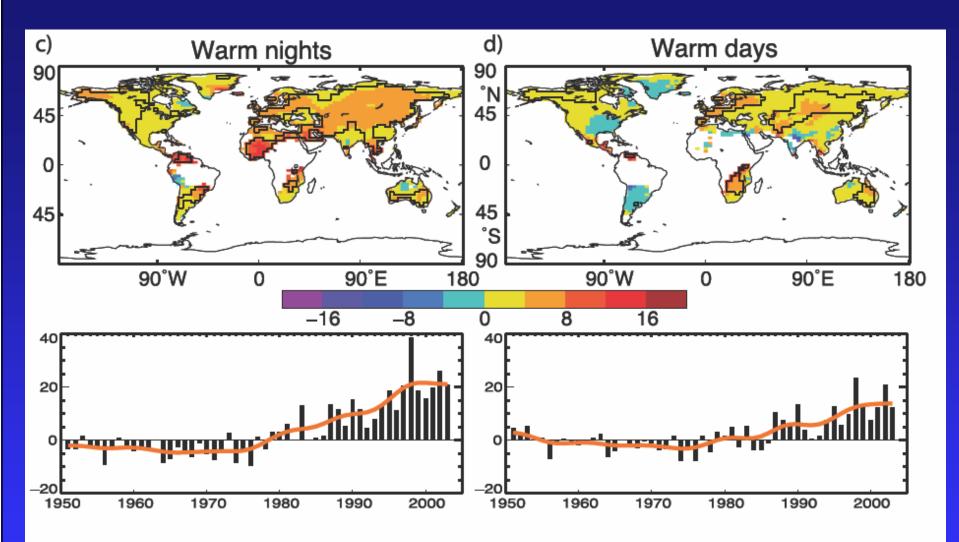
Has there been a Change in Extreme Events like Heat Waves, Droughts, Floods and Hurricanes?

Since 1950, the number of heat waves has increased and widespread increases have occurred in the numbers of warm nights. The extent of regions affected by droughts has also increased as precipitation over land has marginally decreased while evaporation has increased due to warmer conditions. Generally, numbers of heavy daily precipitation events that lead to flooding have increased, but not everywhere. Tropical storm and hurricane frequencies vary considerably from year to year, but evidence suggests substantial increases in intensity and duration since the 1970s. In the extratropics, variations in tracks and intensity of storms reflect variations in major features of the atmospheric circulation, such as the North Atlantic Oscillation.

Table 3.6. Global trends in extremes of temperature and precipitation as measured by the 10th and 90th percentiles (for 1961–1990). Trends with 5 and 95% confidence intervals and levels of significance (**bold**: <1%) were estimated by REML (see Appendix 3.A), which allows for serial correlation in the residuals of the data about the linear trend. All trends are based on annual averages. Values are % per decade. Based on Alexander et al. (2006).

Series	Trend (% pe	
TN10: % incidence of T _{min} below coldest decile.	-1.17 ± 0.20	-1.24 ± 0.44
TN90: % incidence of T _{min} above warmest decile.	1.43 ± 0.42	2.60 ± 0.81
TX10: % incidence of T _{max} below coldest decile.	-0.63 ± 0.16	-0.91 ± 0.48
TX90: % incidence of T _{max} above warmest decile.	0.71 ± 0.35	1.74 ± 0.72
PREC: % contribution of very wet days (above the 95th percentile) to the annual precipitation total.	0.21 ± 0.10	0.41 ± 0.38





FAQ 3.3, Figure 1. Observed trends (days per decade) for 1951 to 2003 in the frequency of extreme temperatures, defined based on 1961 to 1990 values, as maps for the 10th percentile: (a) cold nights and (b) cold days; and 90th percentile: (c) warm nights and (d) warm days. Trends were calculated only for grid boxes that had at least 40 years of data during this period and had data until at least 1999. Black lines enclose regions where trends are significant at the 5% level. Below each map are the global annual time series of anomalies (with respect to 1961 to 1990). The red line shows decadal variations. Trends are significant at the 5% level for all the global indices shown. Adapted from Alexander et al. (2006).

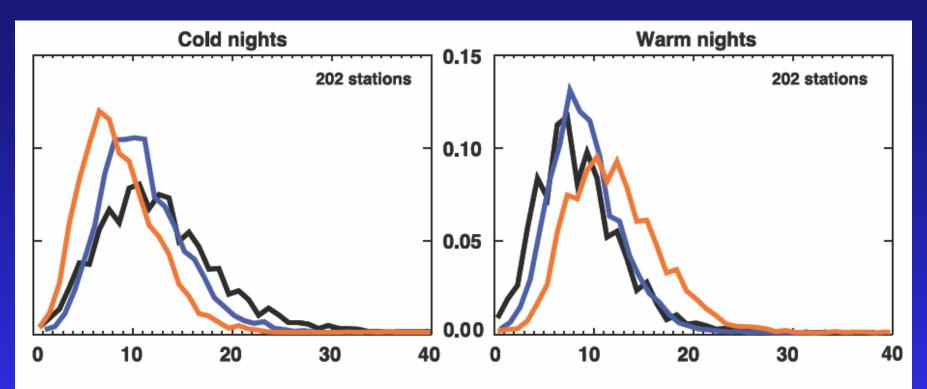
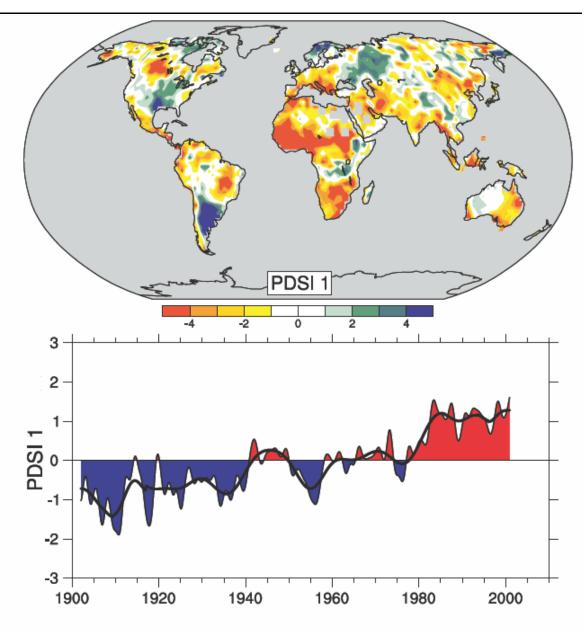
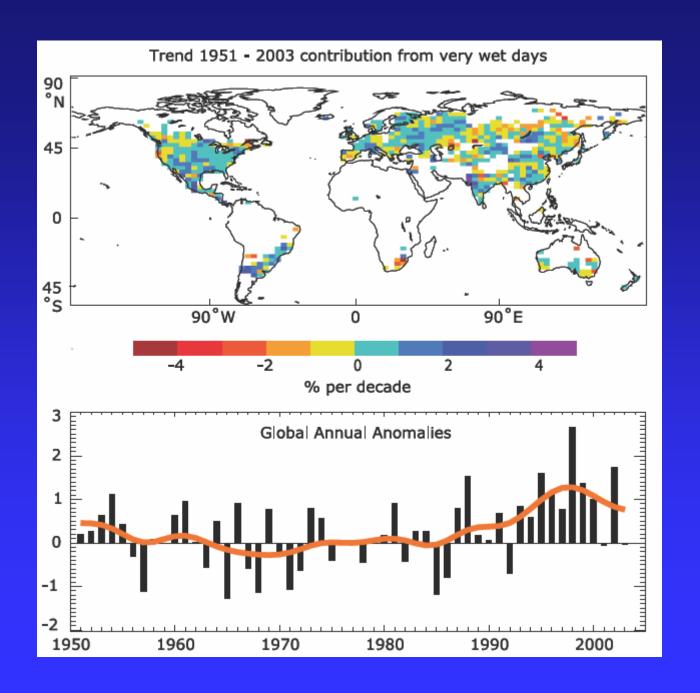


Figure 3.38. Annual probability distribution functions for temperature indices for 202 global stations with at least 80% complete data between 1901 and 2003 for three time periods: 1901 to 1950 (black), 1951 to 1978 (blue) and 1979 to 2003 (red). The x-axis represents the percentage of time during the year when the indicators were below the 10th percentile for cold nights (left) or above the 90th percentile for warm nights (right). From Alexander et al. (2006).



FAQ 3.2, Figure 1. The most important spatial pattern (top) of the monthly Palmer Drought Severity Index (PDSI) for 1900 to 2002. The PDSI is a prominent index of drought and measures the cumulative deficit (relative to local mean conditions) in surface land moisture by incorporating previous precipitation and estimates of moisture drawn into the atmosphere (based on atmospheric temperatures) into a hydrological accounting system. The lower panel shows how the sign and strength of this pattern has changed since 1900. Red and orange areas are drier (wetter) than average and blue and green areas are wetter (drier) than average when the values shown in the lower plot are positive (negative). The smooth black curve shows decadal variations. The time series approximately corresponds to a trend, and this pattern and its variations account for 67% of the linear trend of PDSI from 1900 to 2002 over the global land area. It therefore features widespread increasing African drought, especially in the Sahel, for instance. Note also the wetter areas, especially in eastern North and South America and northern Eurasia. Adapted from Dai et al. (2004b).



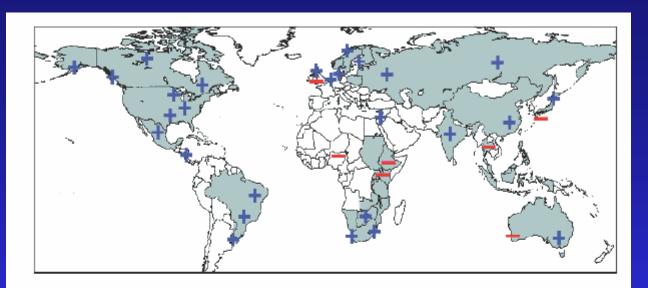


Figure 3.39. (Top) Observed trends (% per decade) for 1951 to 2003 in the contribution to total annual precipitation from very wet days (95th percentile). Trends were only calculated for grid boxes where both the total and the 95th percentile had at least 40 years of data during this period and had data until at least 1999. (Middle) Anomalies (%) of the global annual time series (with respect to 1961 to 1990) defined as the percentage change of contributions of very wet days from the base period average (22.5%). The smooth red curve shows decadal variations (see Appendix 3.A). From Alexander et al. (2006). (Bottom) Regions where disproportionate changes in heavy and very heavy precipitation during the past decades were documented as either an increase (+) or decrease (-) compared to the change in the annual and/or seasonal precipitation (updated from Groisman et al., 2005). Thresholds used to define "heavy" and "very heavy" precipitation vary by season and region. However, changes in heavy precipitation frequencies are always greater than changes in precipitation totals and, in some regions, an increase in heavy and/or very heavy precipitation occurred while no change or even a decrease in precipitation totals was observed.

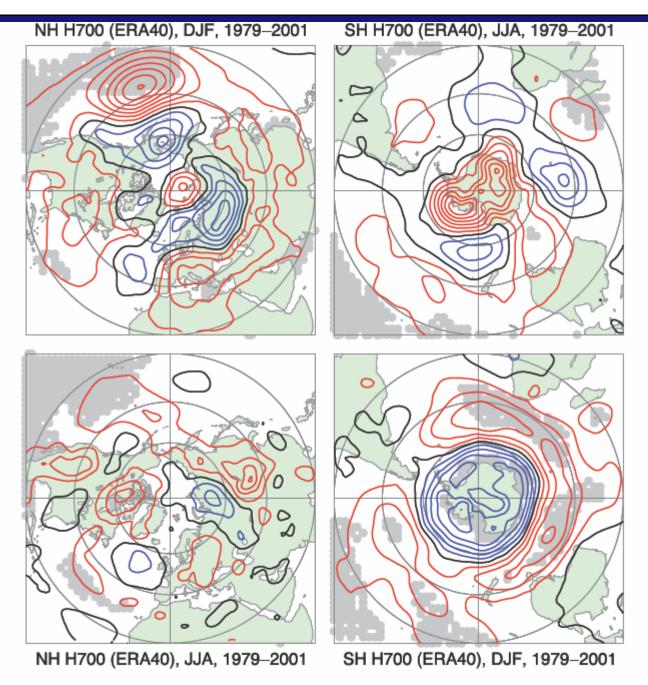


Figure 3.24. Linear trends in ERA-40 700 hPa geopotential height from 1979 to 2001 for DJF (top left and bottom right) and JJA (bottom left and top right), for the NH (left) and SH (right). Trends are contoured in 5 gpm per decade and are calculated from seasonal means of daily 1200 UTC fields. Red contours are positive, blue negative and black zero; the grey background indicates 1% statistical significance using a standard least squares F-test and assuming independence between years.

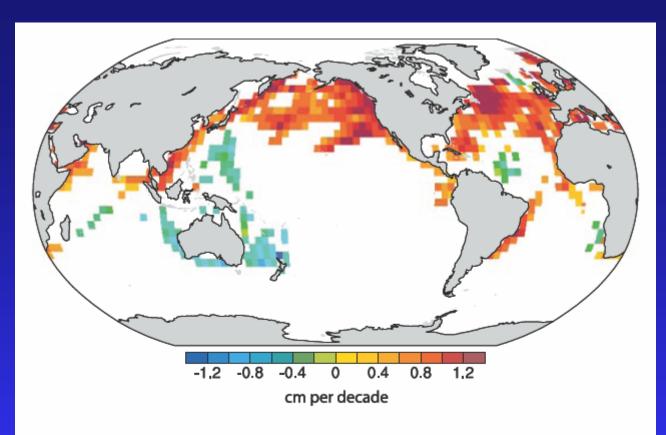


Figure 3.25. Estimates of linear trends in significant wave height (cm per decade) for regions along the major ship routes of the global ocean for 1950 to 2002. Trends are shown only for locations where they are significant at the 5% level. Adapted from Gulev and Grigorieva (2004).

Table 3.7. Definition of phenomena used to assess extremes in Table 3.8

PHENOMENON	Definition
Low-temperature days/ nights and frost days	Percentage of days with temperature (maximum for days, minimum for nights) not exceeding some threshold, either fixed (frost days) or varying regionally (cold days/cold nights), based on the 10th percentile of the daily distribution in the reference period (1961–1990).
High-temperature days/nights	See low-temperature days/nights, but now exceeding the 90th percentile.
Cold spells/snaps	Episode of several consecutive low-temperature days/nights.
Warm spells (heat waves)	Episode of several consecutive high-temperature days/nights.
Cool seasons/warm seasons	Seasonal averages (rather than daily temperatures) exceeding some threshold.
Heavy precipitation events (events that occur every year)	Percentage of days (or daily precipitation amount) with precipitation exceeding some threshold, either fixed or varying regionally, based on the 95th or 99th percentile of the daily distribution in the reference period (1961–1990).
Rare precipitation events (with return periods >~10 yr)	As for heavy precipitation events, but for extremes further into the tail of the distribution.
Drought (season/year)	Precipitation deficit; or based on the PDSI (see Box 3.1).

Tropical cyclones (frequency, intensity, track, peak wind, peak precipitation)

Extreme extratropical storms (frequency, intensity, track, surface wind, wave height)

Small-scale severe weather phenomena

Tropical storm with thresholds crossed in terms of estimated wind speed and organisation. Hurricanes in categories 1 to 5, according to the Saffir-Simpson scale, are defined as storms with wind speeds of 33 to 42 m s⁻¹, 43 to 49 m s⁻¹, 50 to 58 m s⁻¹, 59 to 69 m s⁻¹, and >70 m s⁻¹, respectively. NOAA's ACE index is a measure of the total seasonal activity that accounts for the collective intensity and duration of tropical storms and hurricanes during a given tropical cyclone season.

Intense low-pressure systems that occur throughout the mid-latitudes of both hemispheres fueled by temperature gradients and acting to reduce them.

Extreme events, such as tornadoes, hail, thunderstorms, dust storms and other severe local weather.

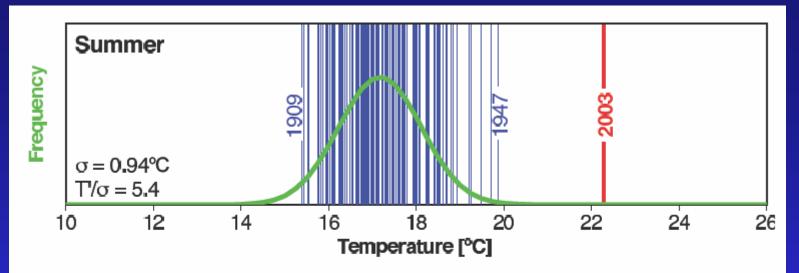
Table 3.8. Change in extremes for phenomena over the specified region and period, with the level of confidence and section where the phenomenon is discuss

PHENOMENON	Change	Region	Period	Confidence
Low-temperature days/nights and frost days	Decrease, more so for nights than days	Over 70% of global land area	1951–2003 (last 150 years for Europe and China)	Very likely
High-temperature days/nights	Increase, more so for nights than days	Over 70% of global land area	1951–2003	Very likely
Cold spells/snaps (episodes of several days)	Insufficient studies, but daily temperature changes imply a decrease			
Warm spells (heat waves) (episodes of several days)	Increase: implicit evidence from changes of daily temperatures	Global	1951–2003	Likely
Cool seasons/ warm seasons (seasonal averages)	Some new evidence for changes in inter-seasonal variability	Central Europe	1961–2004	Likely

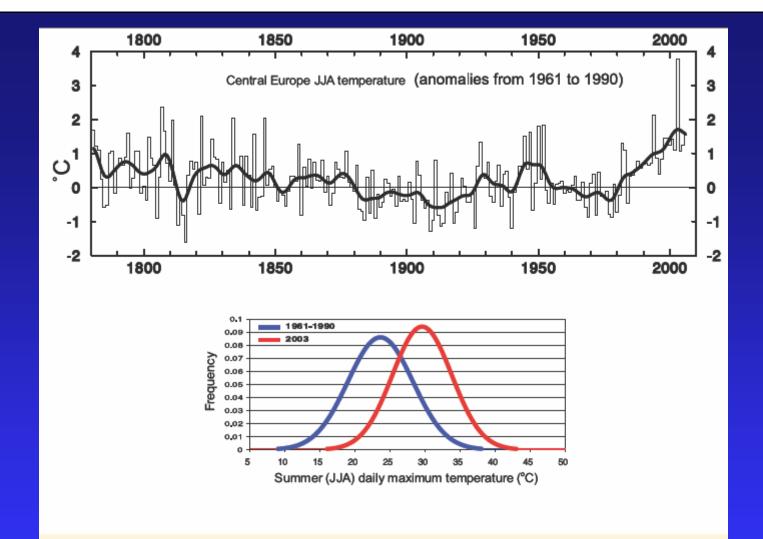
Heavy precipitation events (that occur every year)	Increase, generally beyond that expected from changes in the mean (disproportionate)	Many mid-latitude regions (even where reduction in total precipitation)	1951–2003	Likely
Rare precipitation events (with return periods > ~10 yr)	Increase	Only a few regions have sufficient data for reliable trends (e.g., UK and USA)	Various since 1893	Likely (consistent with changes inferred for more robust statistics)
Drought (season/year)	Increase in total area affected	Many land regions of the world	Since 1970s	Likely
Tropical cyclones	Trends towards longer lifetimes and greater storm intensity, but no trend in frequency	Tropics	Since 1970s	Likely; more confidence in frequency and intensity
Extreme extratropical storms	Net increase in frequency/intensity and poleward shift in track	NH land	Since about 1950	Likely
Small-scale severe weather phenomena	Insufficient studies for assessment			

Frequently Asked Question 9.1 Can Individual Extreme Events be Explained by Greenhouse Warming?

Changes in climate extremes are expected as the climate warms in response to increasing atmospheric greenhouse gases resulting from human activities, such as the use of fossil fuels. However, determining whether a specific, single extreme event is due to a specific cause, such as increasing greenhouse gases, is difficult, if not impossible, for two reasons: 1) extreme events are usually caused by a combination of factors and 2) a wide range of extreme events is a normal occurrence even in an unchanging climate. Nevertheless, analysis of the warming observed over the past century suggests that the likelihood of some extreme events, such as heat waves, has increased due to greenhouse warming, and that the likelihood of others, such as frost or extremely cold nights, has decreased. For example, a recent study estimates that human influences have more than doubled the risk of a very hot European summer like that of 2003.



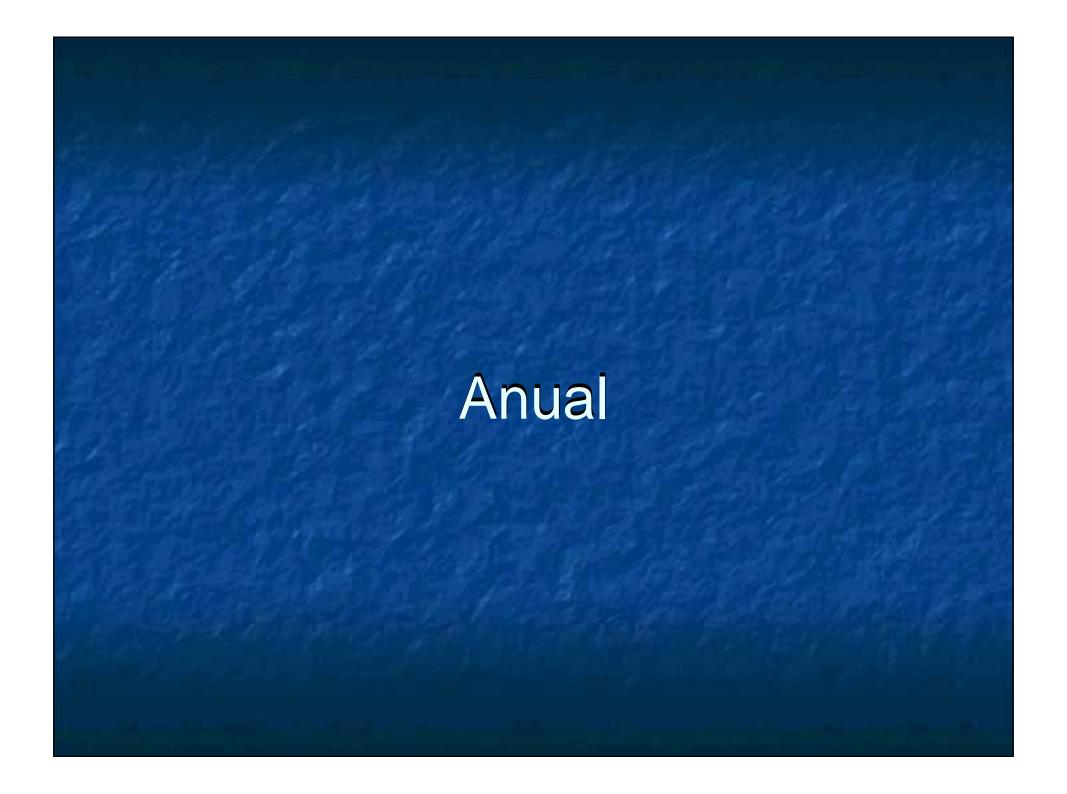
FAQ 9.1, Figure 1. Summer temperatures in Switzerland from 1864 to 2003 are, on average, about 17°C, as shown by the green curve. During the extremely hot summer of 2003, average temperatures exceeded 22°C, as indicated by the red bar (a vertical line is shown for each year in the 137-year record). The fitted Gaussian distribution is indicated in green. The years 1909, 1947 and 2003 are labelled because they represent extreme years in the record. The values in the lower left corner indicate the standard deviation (a) and the 2003 anomaly normalised by the 1864 to 2000 standard deviation (T'/a). From Schär et al. (2004).



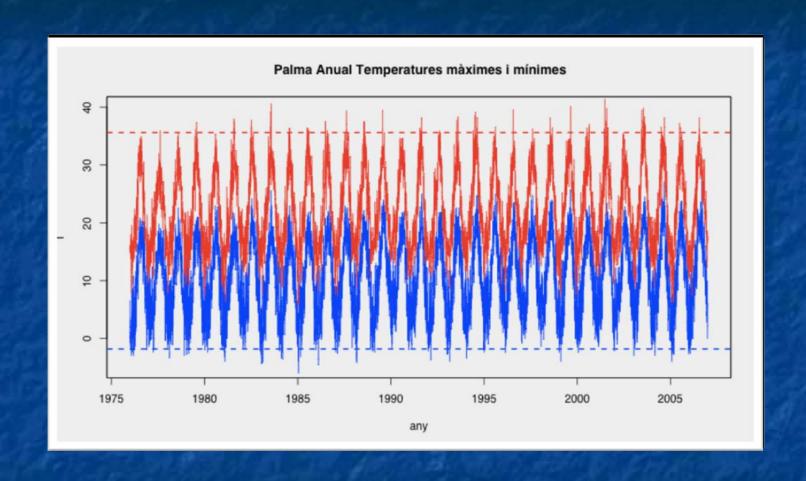
Box 3.6, Figure 2. Long time series of JJA temperature anomalies in Central Europe relative to the 1961 to 1990 mean (top). The smooth curve shows decadal variations (see Appendix 3.A). In the summer of 2003, the value of 3.8°C far exceeded the next largest anomaly of 2.4°C in 1807, and the highly smoothed Gaussian distribution (bottom) of maximum temperatures (red) compared with normal (blue) at Basel, Switzerland (Beniston and Diaz, 2004) shows how the whole distribution shifted.

Extremos de Temperatura

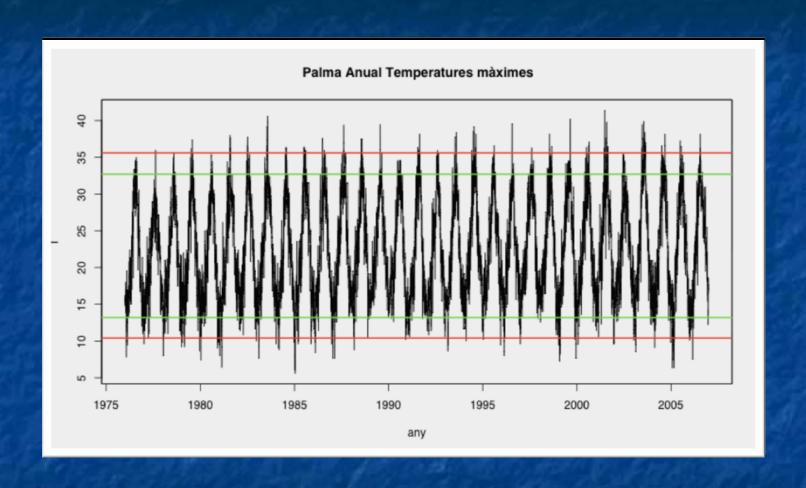
Mallorca



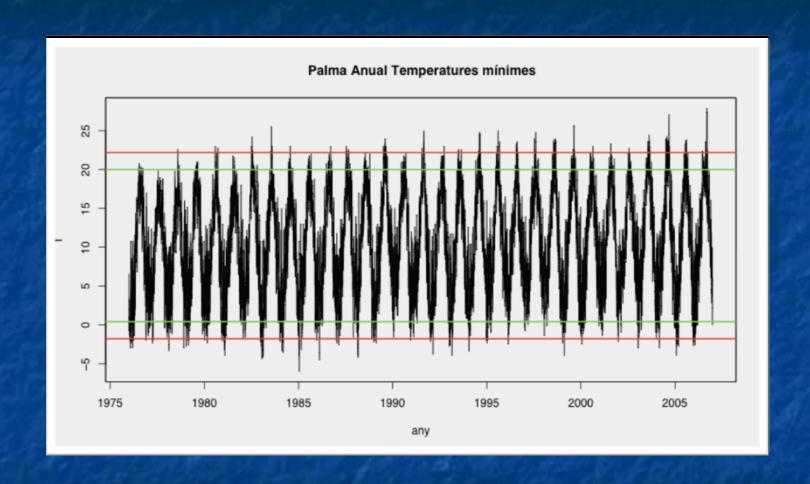
Temperaturas máximas i mínimas diarias



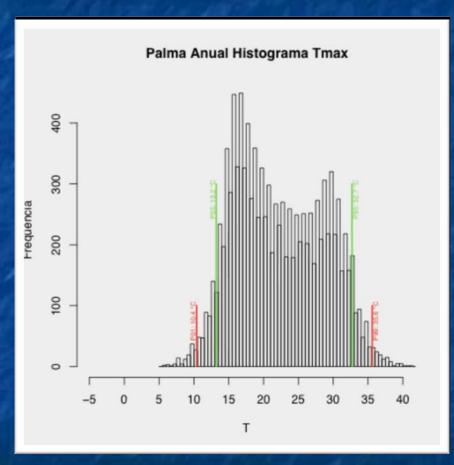
Temperaturas máximas diarias

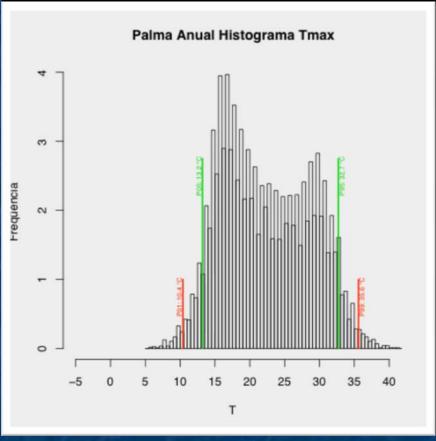


Temperaturas mínimas diarias

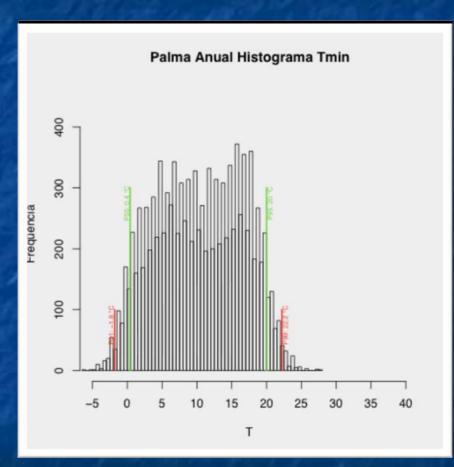


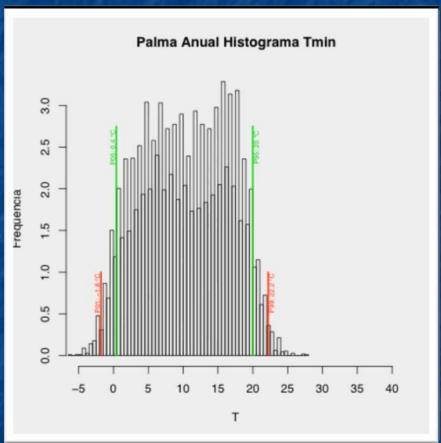
Histograma temperaturas máximas





Histograma temperaturas mínimas



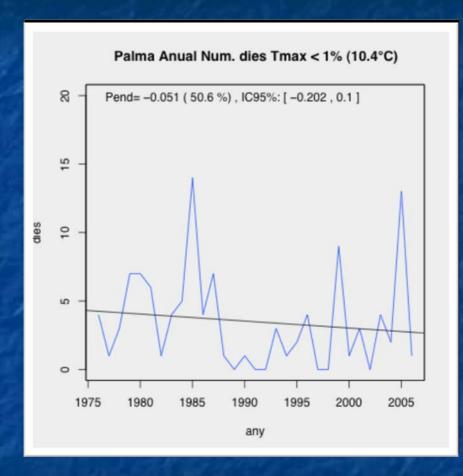


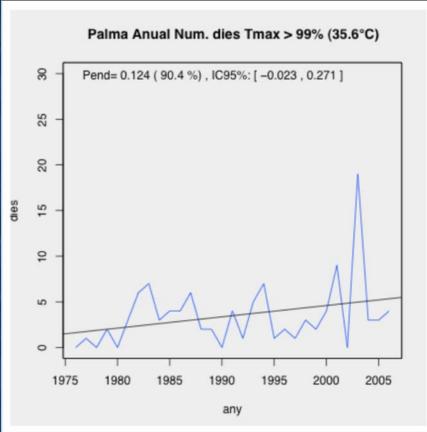
Temperaturas - Extremos

Máximas			
Tmax 1%	10.4°C	Tmax 99%	35.6°C
Tmax 5%	13.2°C	Tmax 95%	32.7°C
Tmax 10%	14.6°C	Tmax 90%	31.2°C
Tmax 20%	16.2°C	Tmax 80%	29.0°C

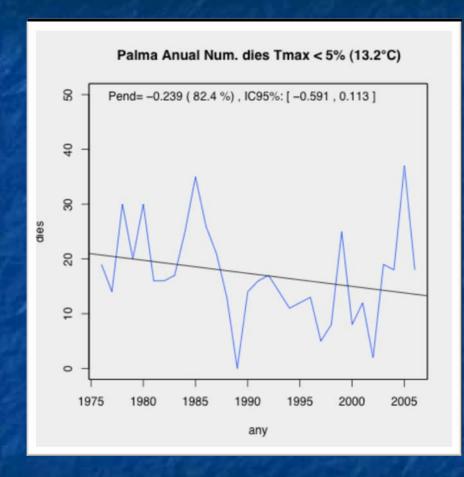
Mínimas				
Tmin 1%	-1.8°C	Tmin 99%	22.2°C	
Tmin 5%	0.4°C	Tmin 95%	20.0°C	
Tmin 10%	1.8°C	Tmin 90%	18.6°C	
Tmin 20%	4.2°C	Tmin 80%	16.6°C	

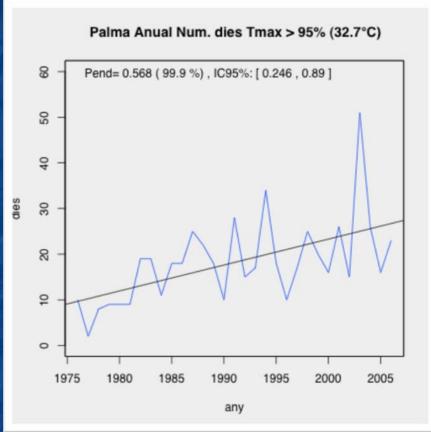
Temperaturas máximas (99%)



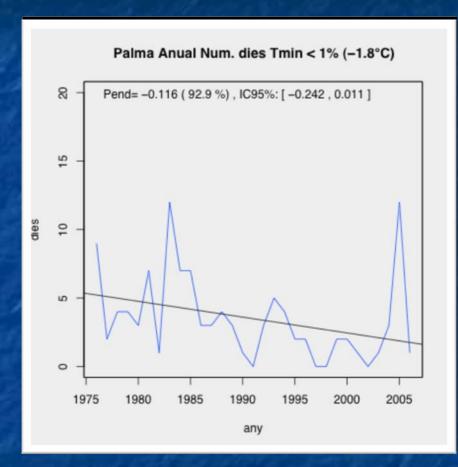


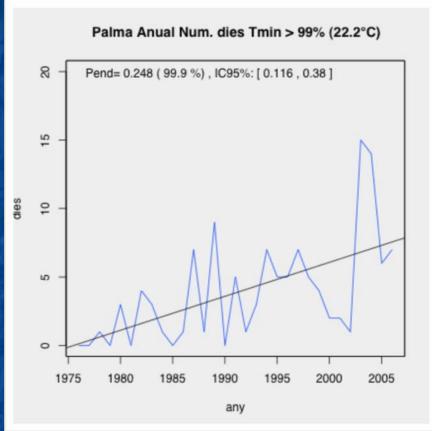
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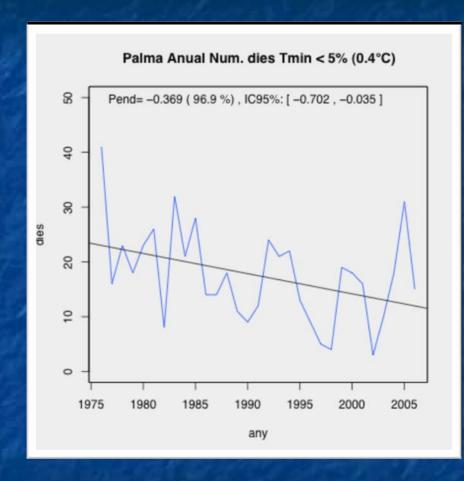


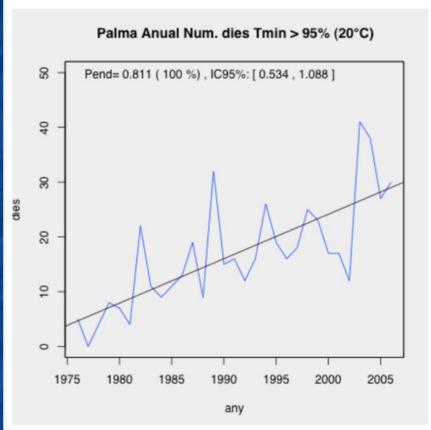
Temperaturas mínimas (99%)





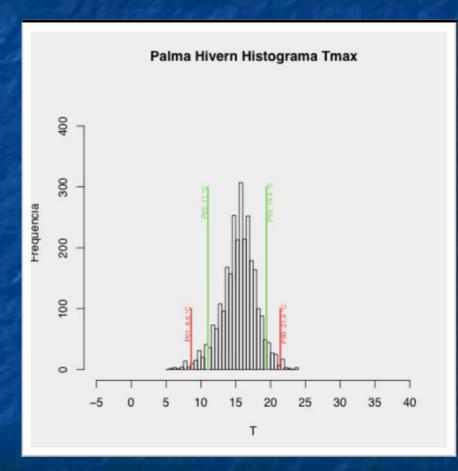
Temperaturas mínimas (95%)

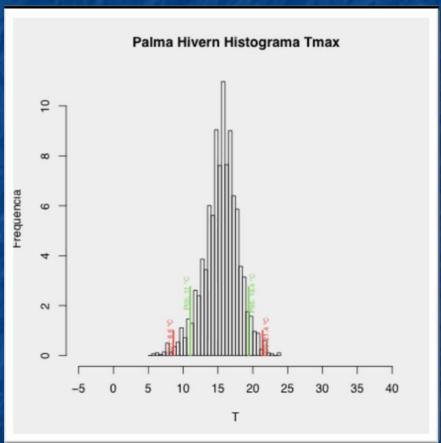




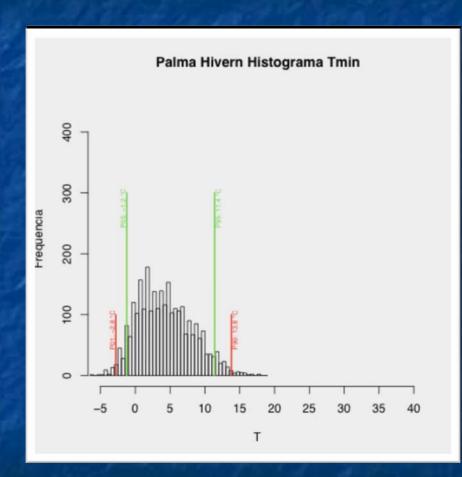


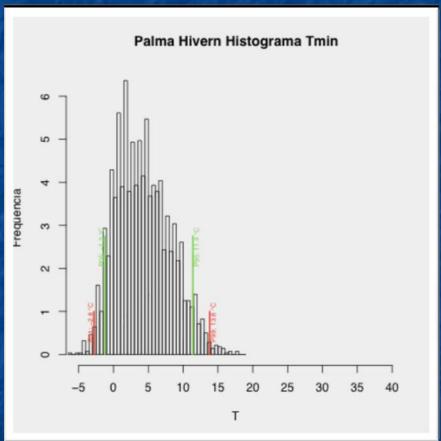
Histograma temperaturas máximas





Histograma temperaturas mínimas



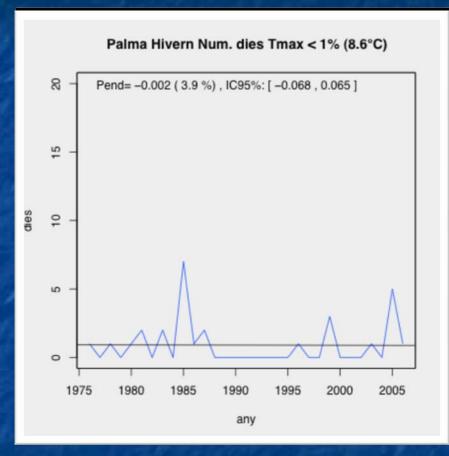


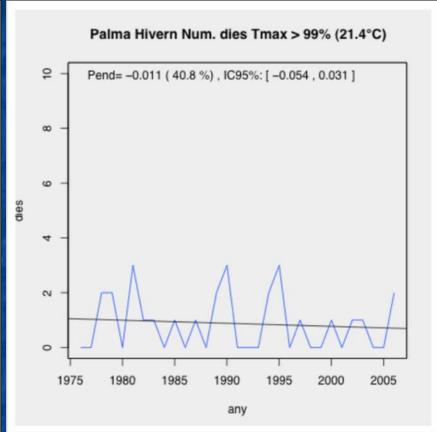
Temperaturas - Extremos

Máximas			
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Tmax 5%	11.0°C	Tmax 95%	19.4°C
Tmax 10%	12.4°C	Tmax 90%	18.5°C
Tmax 20%	13.6°C	Tmax 80%	17.5°C

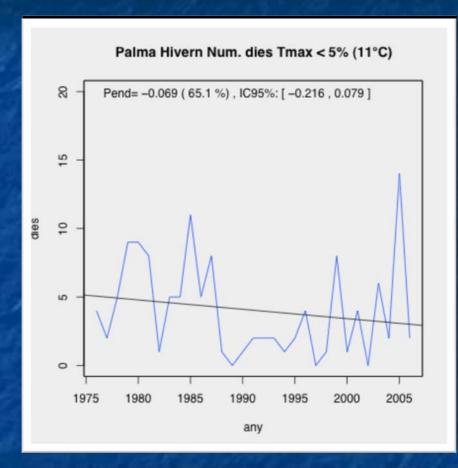
Mínimas				
Tmin 1%	-2.8°C	Tmin 99%	13.8°C	
Tmin 5%	-1.2°C	Tmin 95%	11.4°C	
Tmin 10%	-0.4°C	Tmin 90%	9.6°C	
Tmin 20%	0.8°C	Tmin 80%	7.8°C	

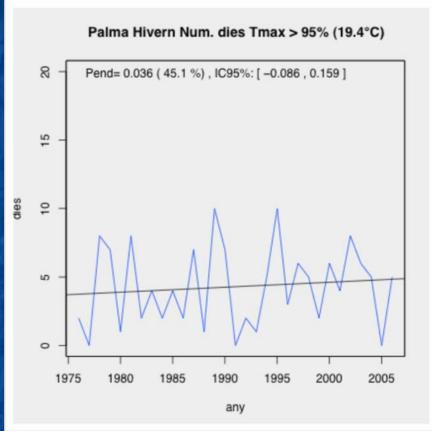
Temperaturas máximas (99%)



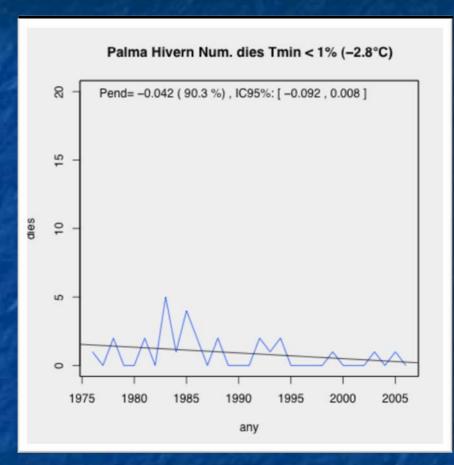


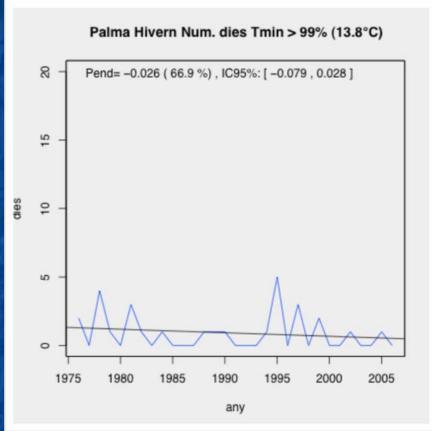
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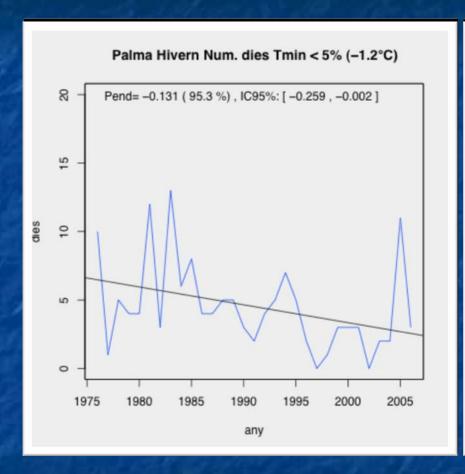


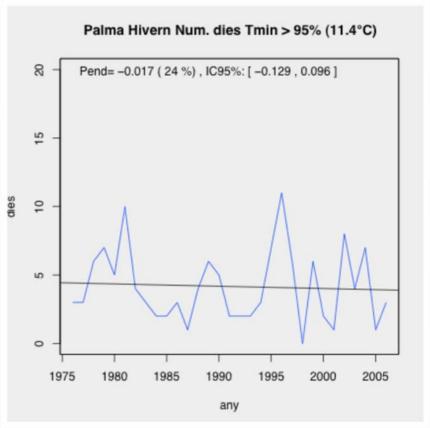
Temperaturas mínimas (99%)

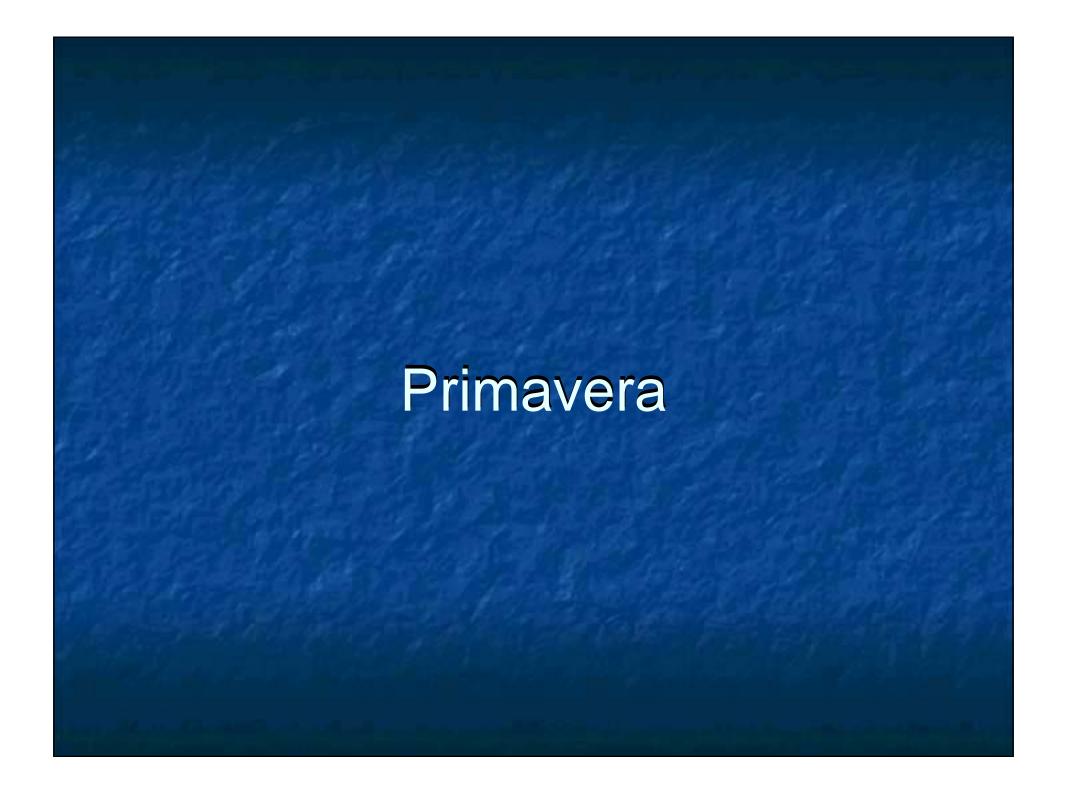




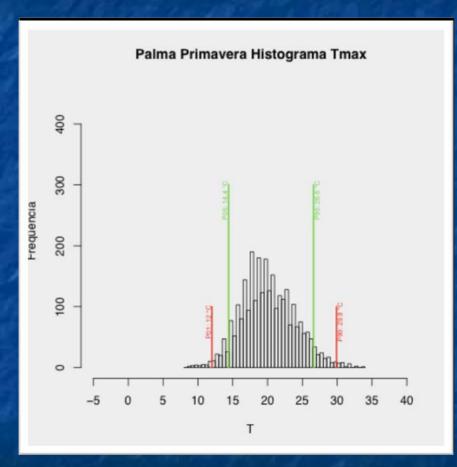
Temperaturas mínimas (95%)

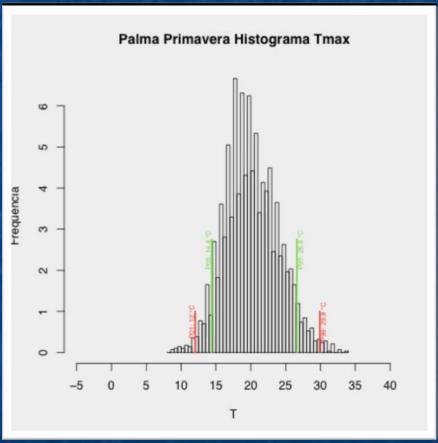




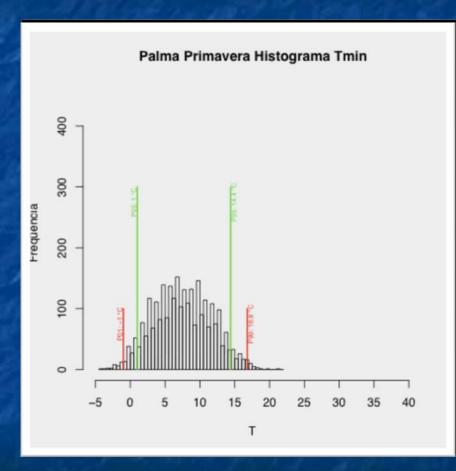


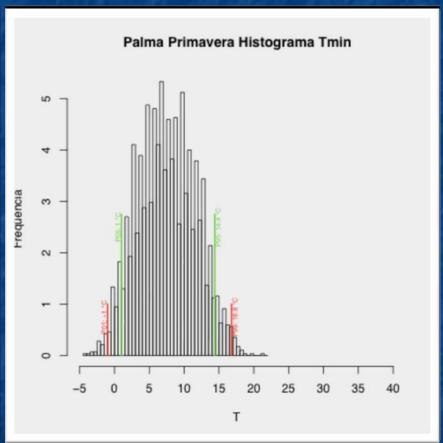
Histograma temperaturas máximas





Histograma temperaturas mínimas



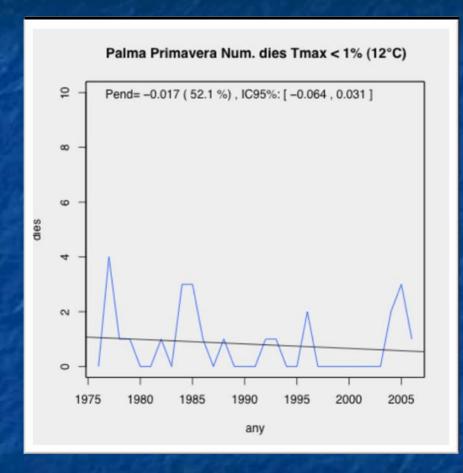


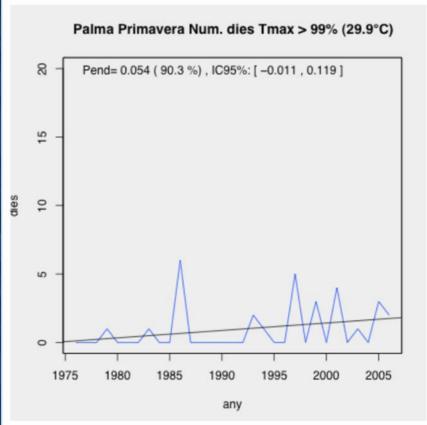
Temperaturas - Extremos

Máximas			
Tmax 1%	12.0°C	Tmax 99%	29.9°C
Tmax 5%	14.4°C	Tmax 95%	26.6°C
Tmax 10%	15.5°C	Tmax 90%	25.4°C
Tmax 20%	17.0°C	Tmax 80%	23.5°C

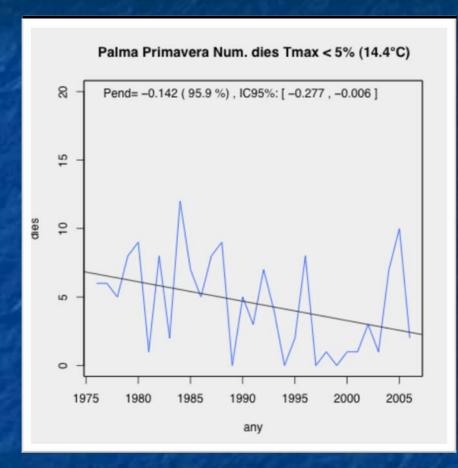
Mínimas			
Tmin 1%	-1.0°C	Tmin 99%	16.8°C
Tmin 5%	1.0°C	Tmin 95%	14.4°C
Tmin 10%	2.2°C	Tmin 90%	13.0°C
Tmin 20%	3.8°C	Tmin 80%	11.4°C

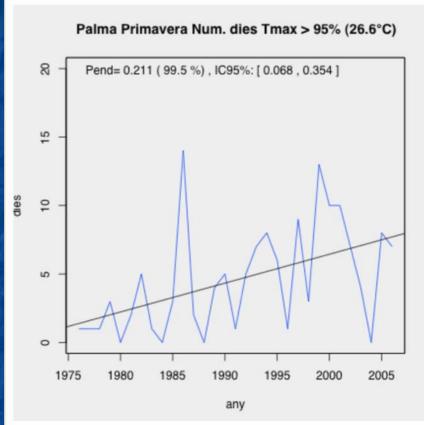
Temperaturas máximas (99%)



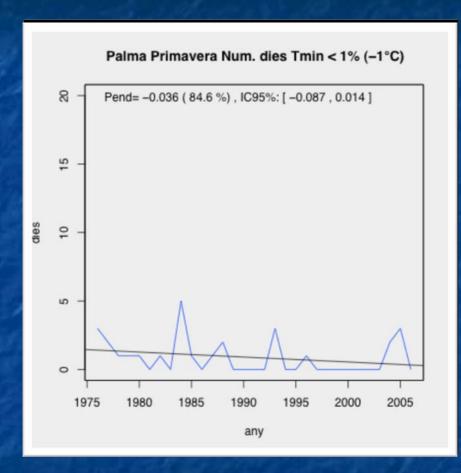


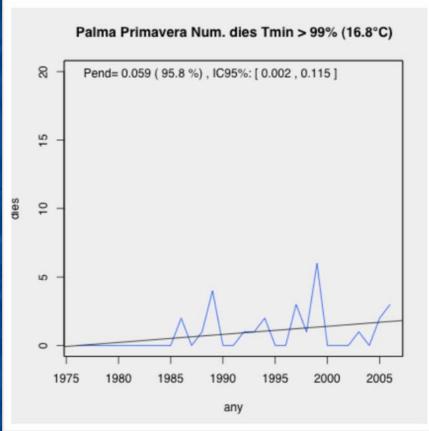
Temperaturas máximas (95%)



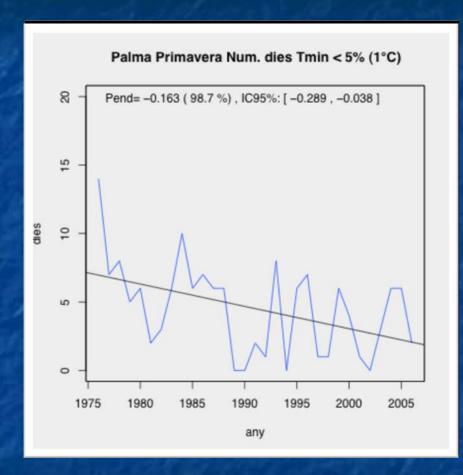


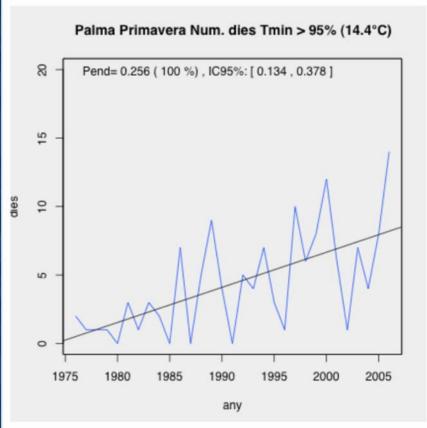
Temperaturas mínimas (99%)





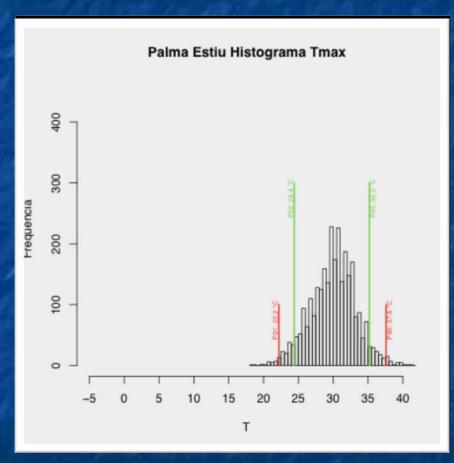
Temperaturas mínimas (95%)

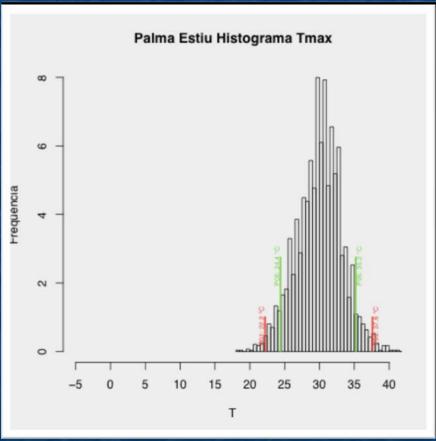




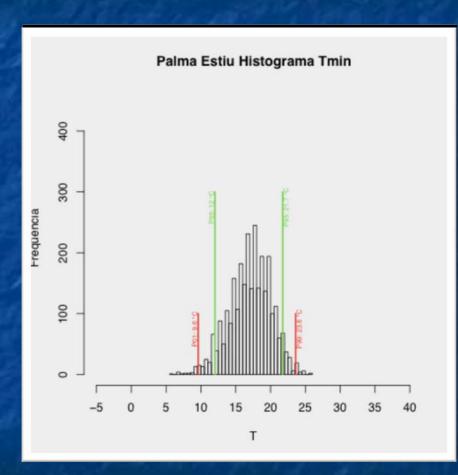


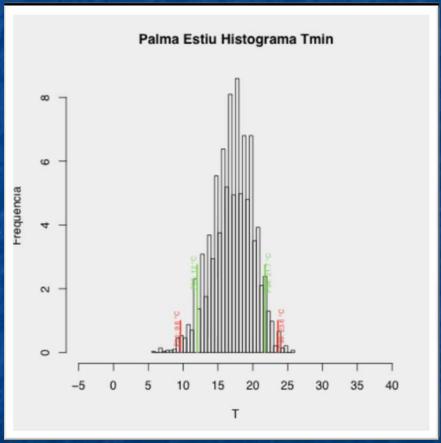
Histograma temperaturas máximas





Histograma temperaturas mínimas



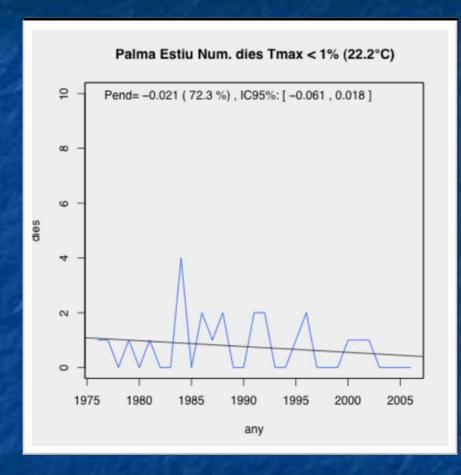


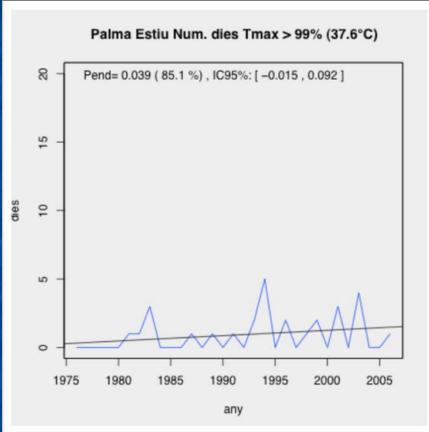
Temperaturas - Extremos

Màximes			
Tmax 1%	22.2°C	Tmax 99%	37.6°C
Tmax 5%	24.4°C	Tmax 95%	35.2°C
Tmax 10%	25.6°C	Tmax 90%	34.0°C
Tmax 20%	27.4°C	Tmax 80%	32.6°C

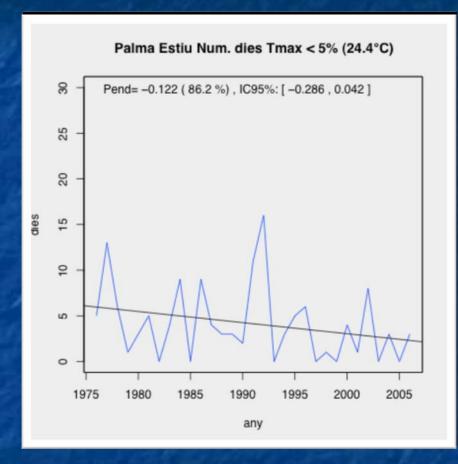
Mínimes			
Tmin 1%	9.6°C	Tmin 99%	23.6°C
Tmin 5%	12.0°C	Tmin 95%	21.7°C
Tmin 10%	13.0°C	Tmin 90%	20.8°C
Tmin 20%	14.6°C	Tmin 80%	19.8°C

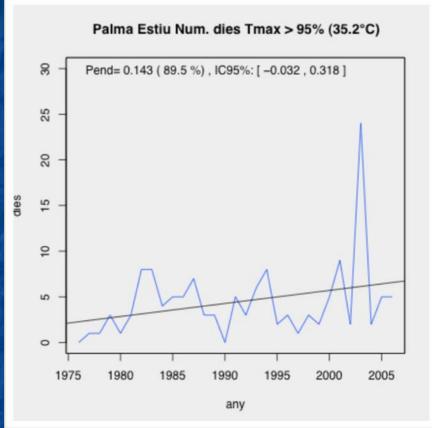
Temperaturas máximas (99%)



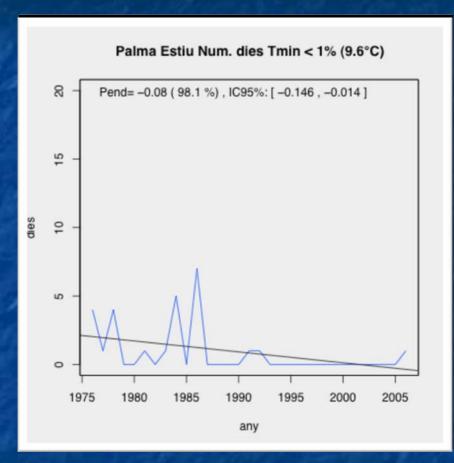


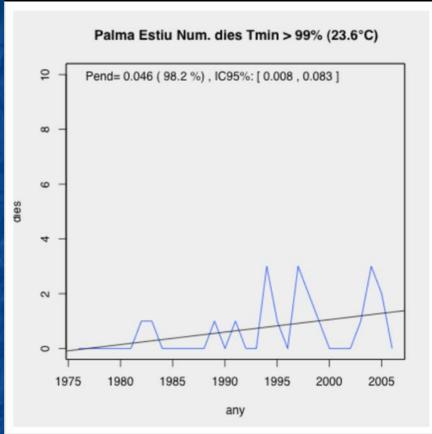
Temperaturas máximas (95%)



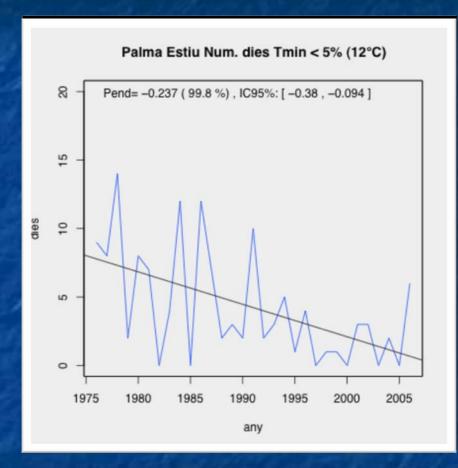


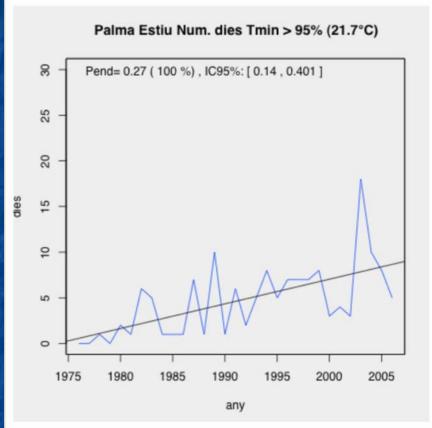
Temperaturas mínimas (99%)

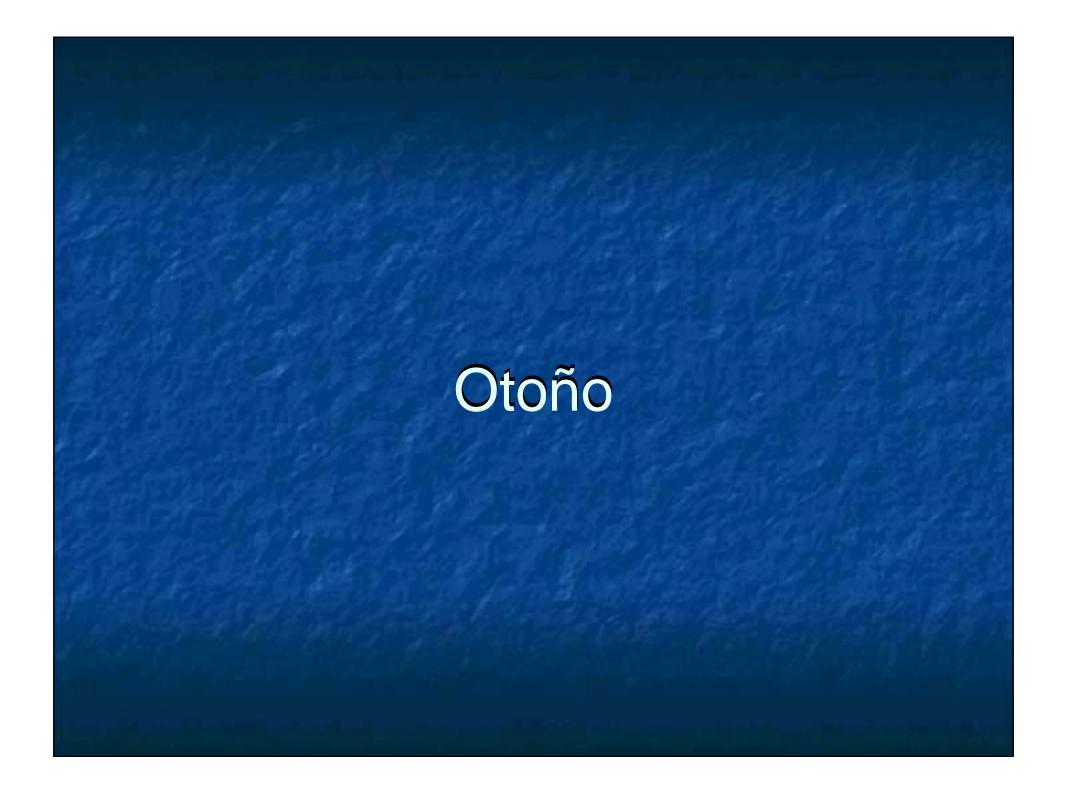




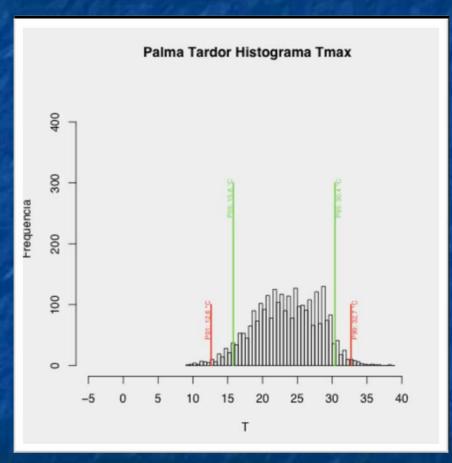
Temperaturas mínimas (95%)

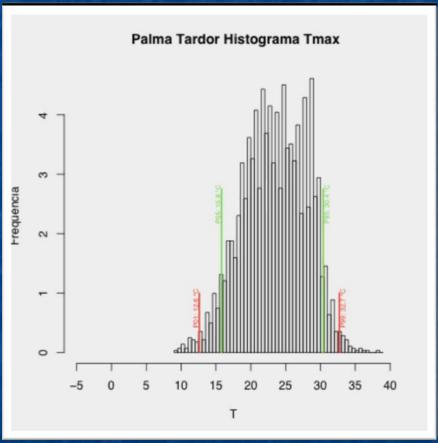




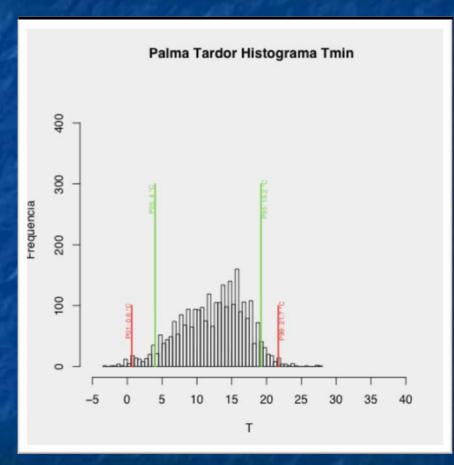


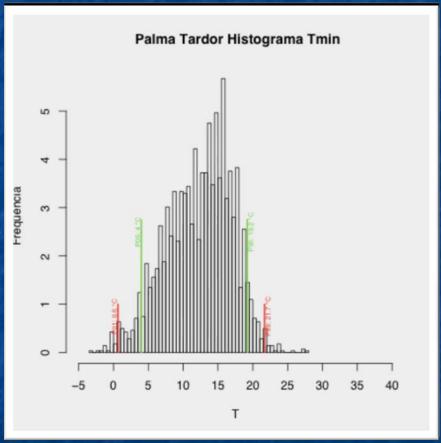
Histograma temperaturas máximas





Histograma temperaturas mínimas



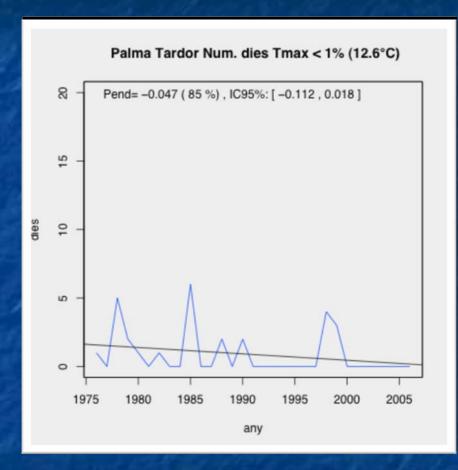


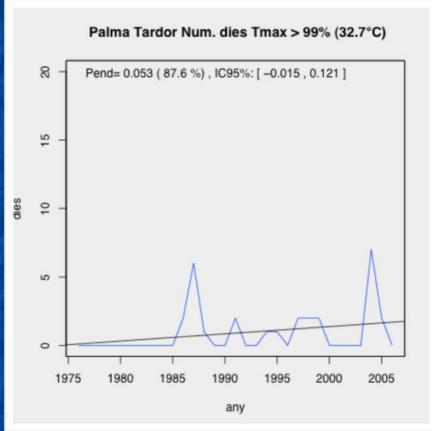
Temperaturas - Extremos

Màximes			
Tmax 1%	12.6°C	Tmax 99%	32.7°C
Tmax 5%	15.8°C	Tmax 95%	30.4°C
Tmax 10%	17.4°C	Tmax 90%	29.4°C
Tmax 20%	19.5°C	Tmax 80%	28.0°C

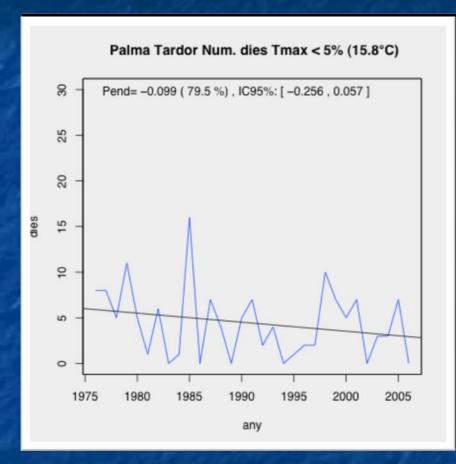
Mínimes			
Tmin 1%	0.6°C	Tmin 99%	21.7°C
Tmin 5%	4.0°C	Tmin 95%	19.2°C
Tmin 10%	5.8°C	Tmin 90%	18.0°C
Tmin 20%	8.2°C	Tmin 80%	16.5°C

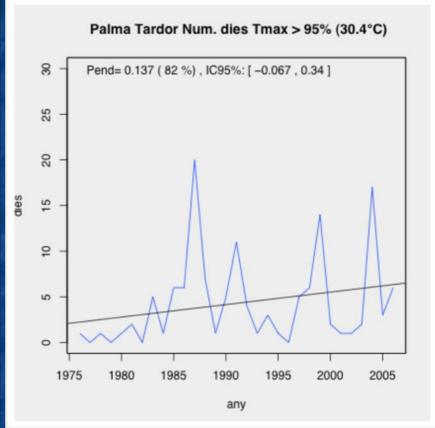
Temperaturas máximas (99%)



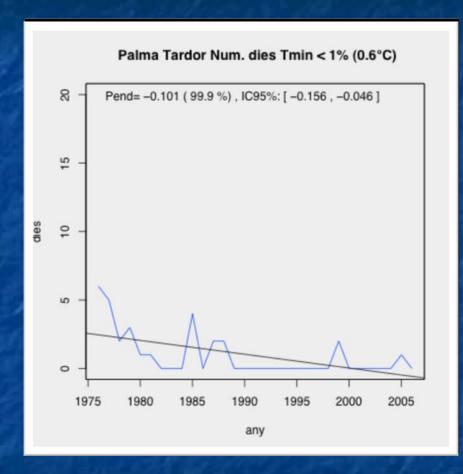


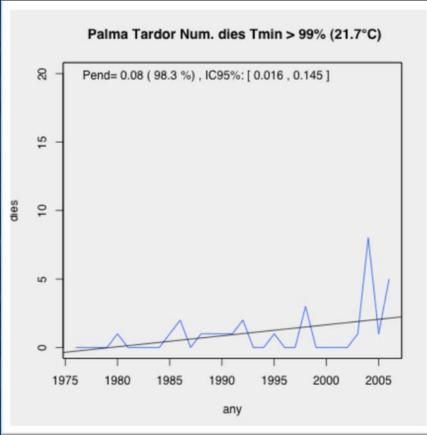
Temperaturas máximas (95%)



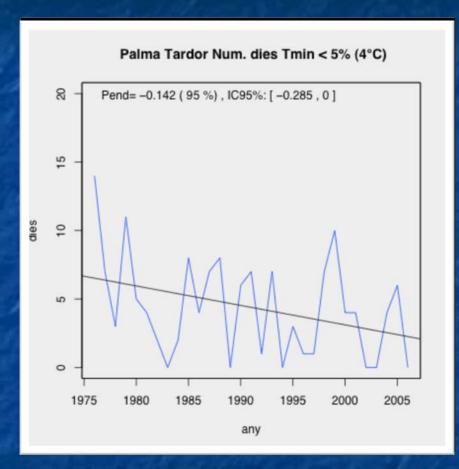


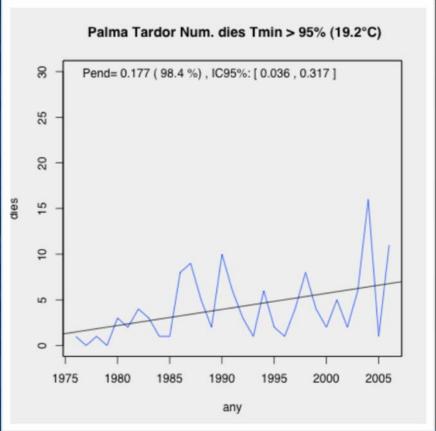
Temperaturas mínimas (99%)





Temperaturas mínimas (95%)





Extremos de Precipitación

Islas Baleares

Daily precipitation categories

Category	Precipitation (mm)
"trace"	<1
a	1-4
b	4-16
c 1	16-32
c2	32-64
d1	64-128
d2	>128

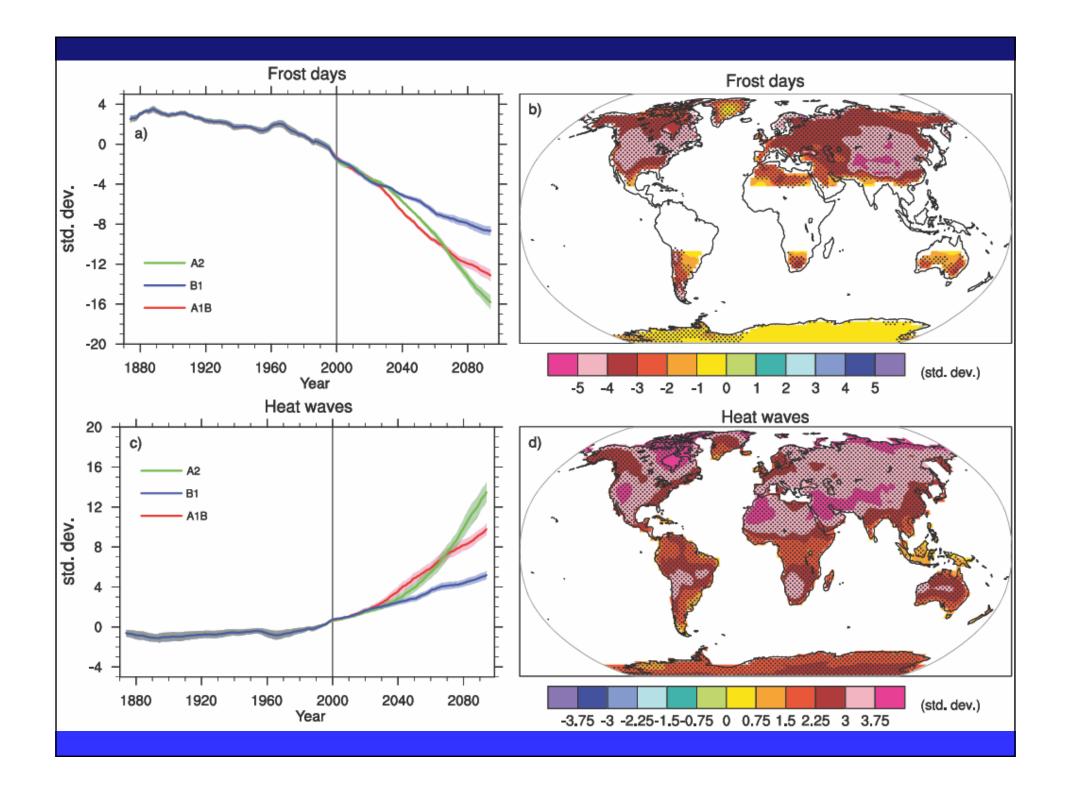
Categories

Category	Mean (days/ yr)	Trend (days/100 yr)	Statistical Confidence in trend (+ or -)
trace	17.65	+3.79	Very likely
а	19.36	+7.01	Virtually certain
b	23.99	-4.86	Likely
c1	7.63	-4.20	Virtually certain
c2	2.79	-1.40	Likely
d1	0.45	+0.07	About as likely as not
d2	0.04	+0.04	About as likely as not

Frequently Asked Question 10.1

Are Extreme Events, Like Heat Waves, Droughts or Floods, Expected to Change as the Earth's Climate Changes?

Yes; the type, frequency and intensity of extreme events are expected to change as Earth's climate changes, and these changes could occur even with relatively small mean climate changes. Changes in some types of extreme events have already been observed, for example, increases in the frequency and intensity of heat waves and heavy precipitation events (see FAQ 3.3).



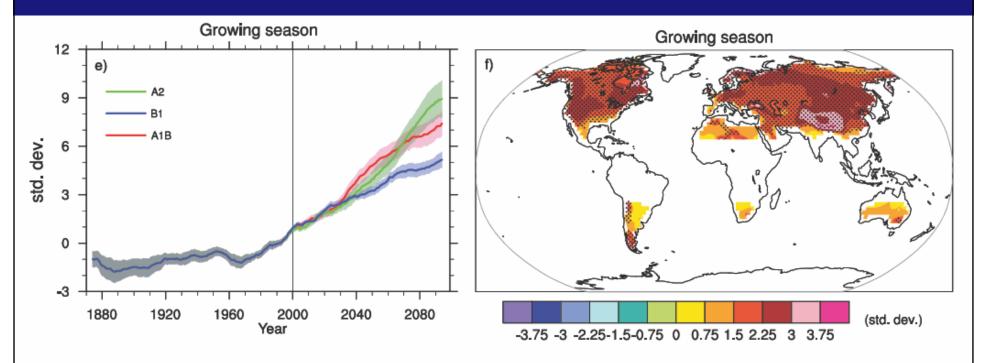
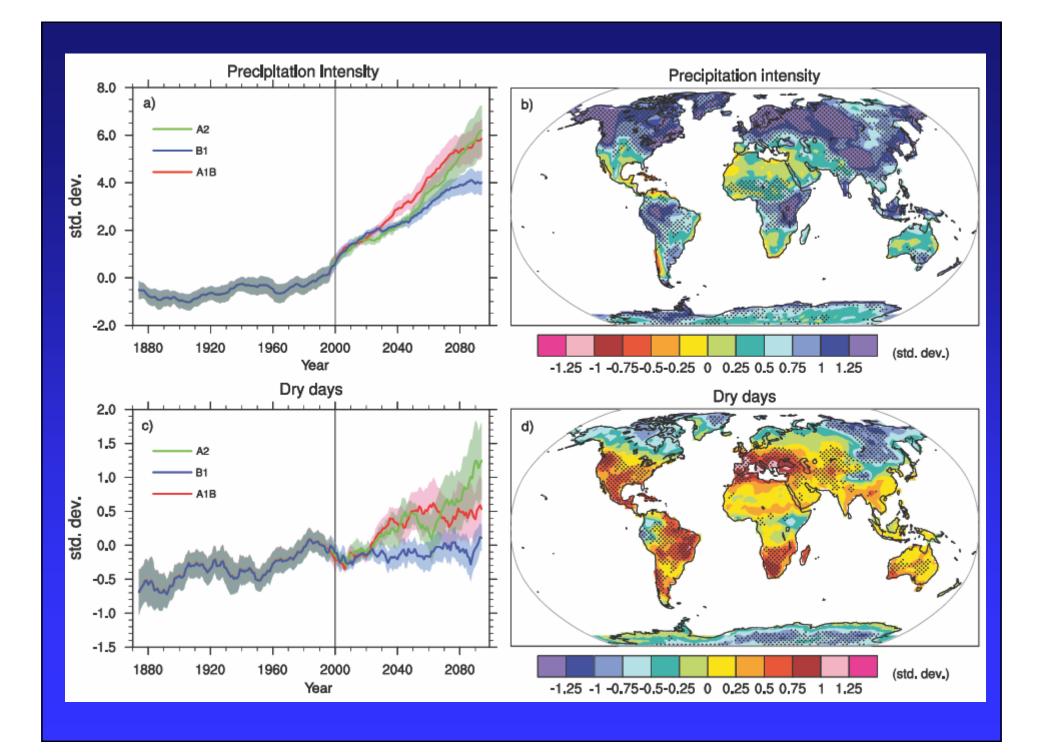


Figure 10.19. Changes in extremes based on multi-model simulations from nine global coupled climate models, adapted from Tebaldi et al. (2006). (a) Globally averaged changes in the frost day index (defined as the total number of days in a year with absolute minimum temperature below 0°C) for a low (SRES B1), middle (SRES A1B) and high (SRES A2) scenario. (b) Changes in spatial patterns of simulated frost days between two 20-year means (2080–2099 minus 1980–1999) for the A1B scenario. (c) Globally averaged changes in heat waves (defined as the longest period in the year of at least five consecutive days with maximum temperature at least 5°C higher than the climatology of the same calendar day). (d) Changes in spatial patterns of simulated heat waves between two 20-year means (2080–2099 minus 1980–1999) for the A1B scenario. (e) Globally averaged changes in growing season length (defined as the length of the period between the first spell of five consecutive days with mean temperature above 5°C and the last such spell of the year). (f) Changes in spatial patterns of simulated growing season length between two 20-year means (2080–2099 minus 1980–1999) for the A1B scenario. Solid lines in (a), (c) and (e) show the 10-year smoothed multi-model ensemble means; the envelope indicates the ensemble mean standard deviation. Stippling in (b), (d) and (f) denotes areas where at least five of the nine models concur in determining that the change is statistically significant. Extreme indices are calculated only over land. Frost days and growing season are only calculated in the extratropics. Extremes indices are calculated following Frich et al. (2002). Each model's time series was centred around its 1980 to 1999 average and normalised (rescaled) by its standard deviation computed (after de-trending) over the period 1960 to 2099. The models were then aggregated into an ensemble average, both at the global and at the grid-box level. Thus, changes are given in units of standard deviations.





HURACANES Y MEDICANES: RESPUESTA AL CALENTAMIENTO GLOBAL

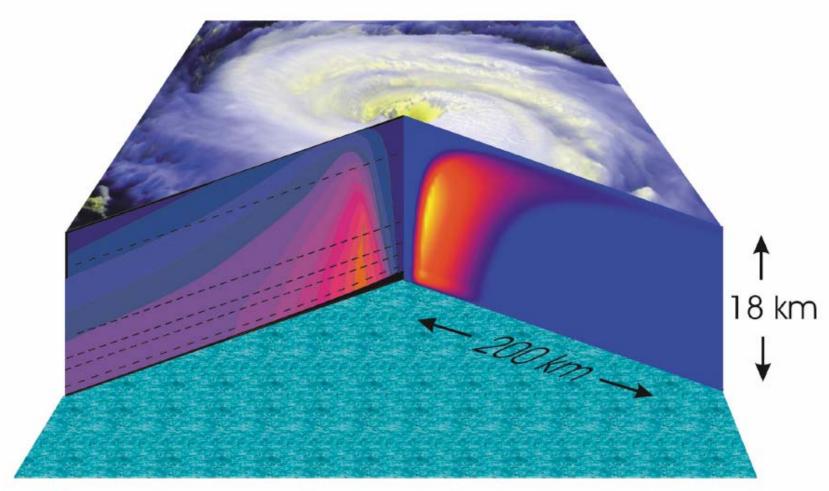


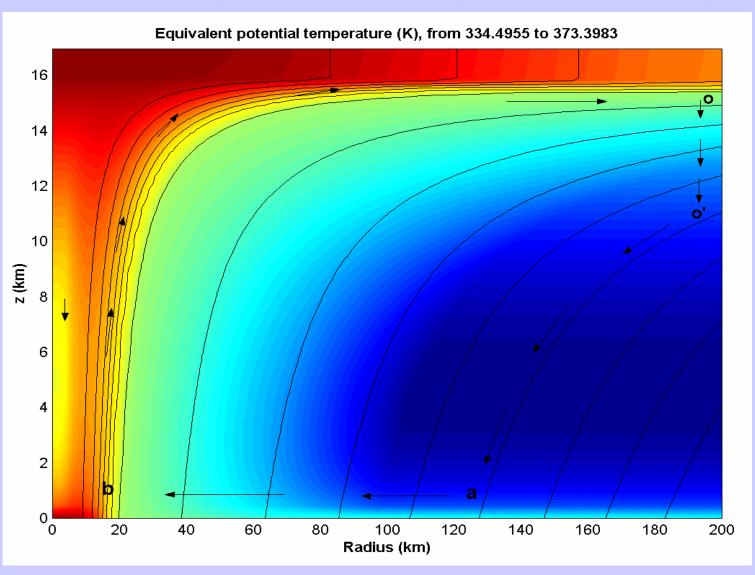
Figure 3 Cutaway view of the structure of a tropical cyclone. The top of the storm is based on a satellite photograph of the cloud structure of Hurricane Fran of 1996. The right-hand cut shows the vertical component of velocity, from a numerical simulation of a hurricane using the model of Emanuel (1995a); maximum values (yellow) are approximately 8 ms⁻¹. The left-hand cut shows the magnitude of the tangential wind component measured in Hurricane Inez of 1966 by aircraft flying at levels indicated by the black dashed lines; from Hawkins & Imbembo (1976). Maximum values are approximately 50 ms⁻¹.

Energía del Huracán (ciclo de Carnot)

a-b: Expansión isotérmica

b-o: Expansión adiabática **o-o'**: Compresión isotérmica

o'-a: Compresión adiabática



Balance Energético (Estado Estacionario)

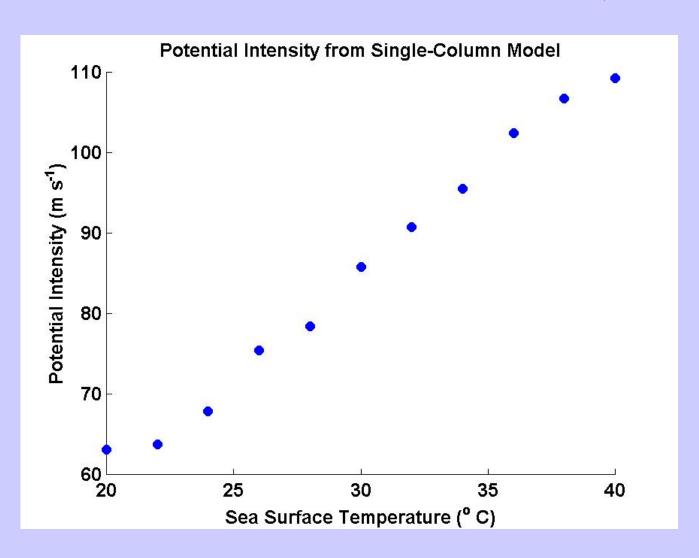
$$\mathsf{P} = 2\pi \frac{T_S - T_O}{T_S} \int_a^b \left[C_k \rho |V| \left(k_0^* - k \right) + C_D \rho |V|^3 \right] r dr$$

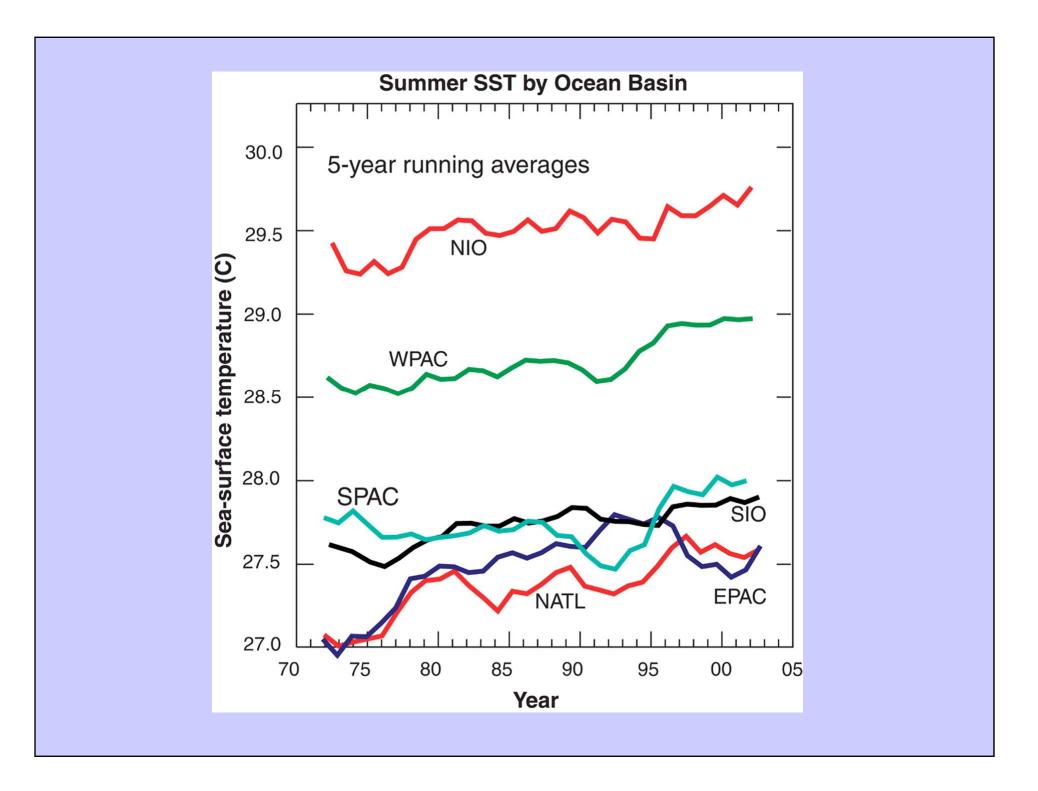
$$D = 2\pi \int_{a}^{b} C_{D} \rho |V|^{3} r dr$$

$$\rightarrow |V_{\text{max}}|^2 \cong \frac{C}{C} \frac{T - T}{C} \left(k_0^* - k\right)$$

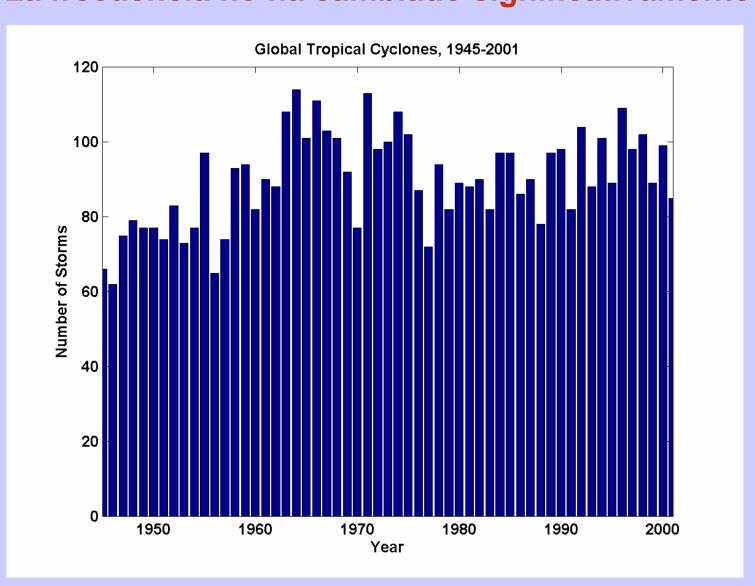
→ P_{min} using the gradient-wind relationship

TEORÍA: La Intensidad Potencial Aumenta con la SST (Emanuel, 1987)

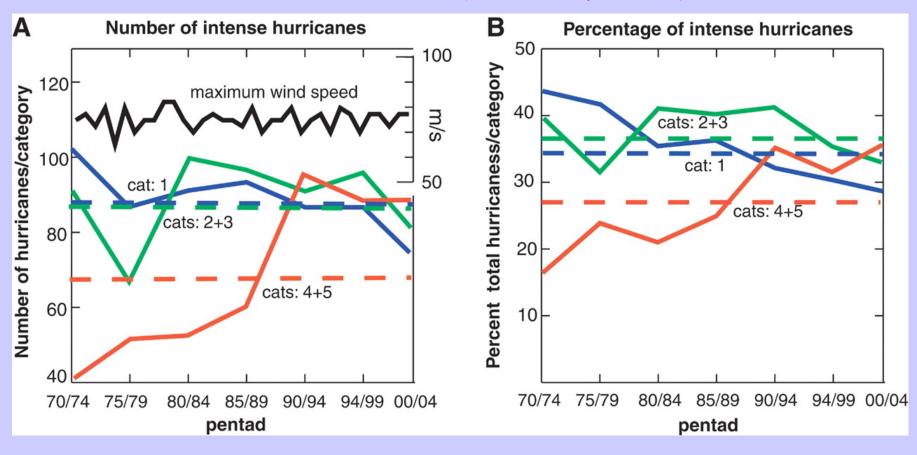




La frecuencia no ha cambiado significativamente



Webster et al. (Science, 2005)



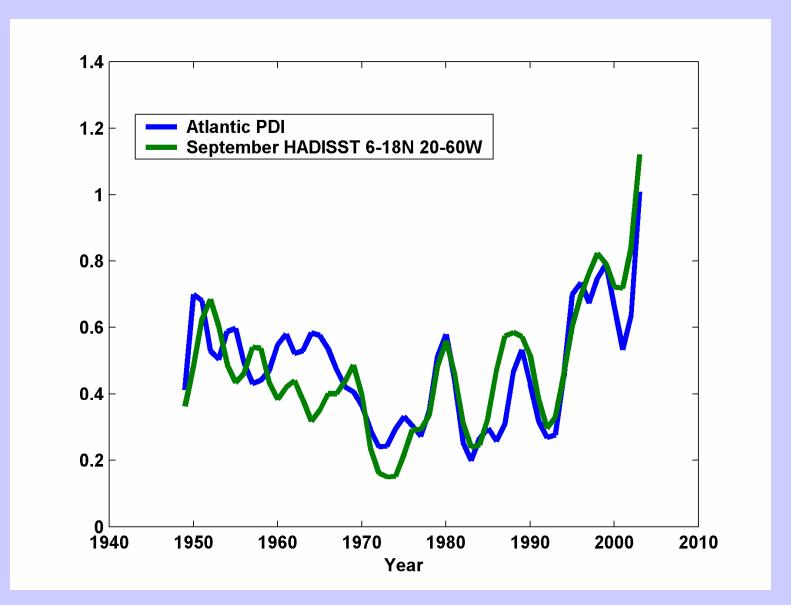
Medida de la actividad del ciclón tropical

Power dissipation =
$$2\pi \int_0^{\tau} \int_0^{r_0} C_D \rho |\mathbf{V}|^3 r dr dt$$
.

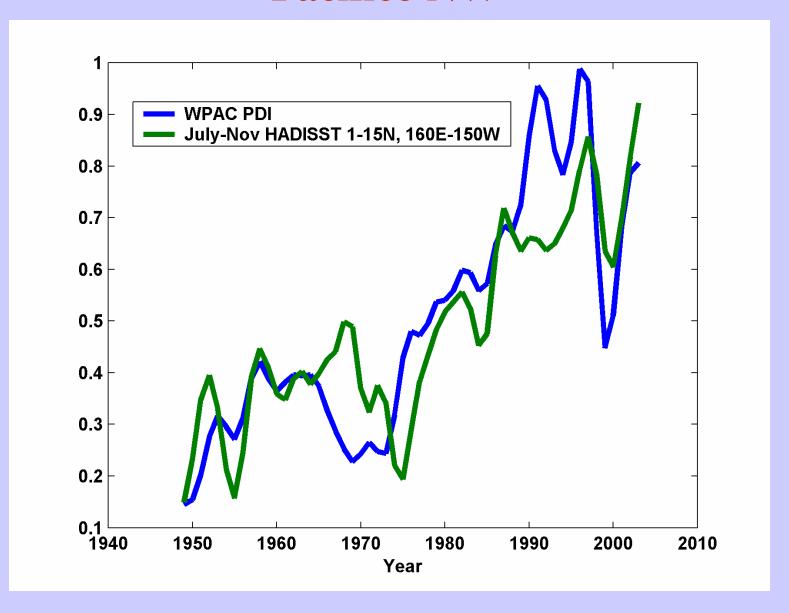
Simplified "Power Dissipation Index":

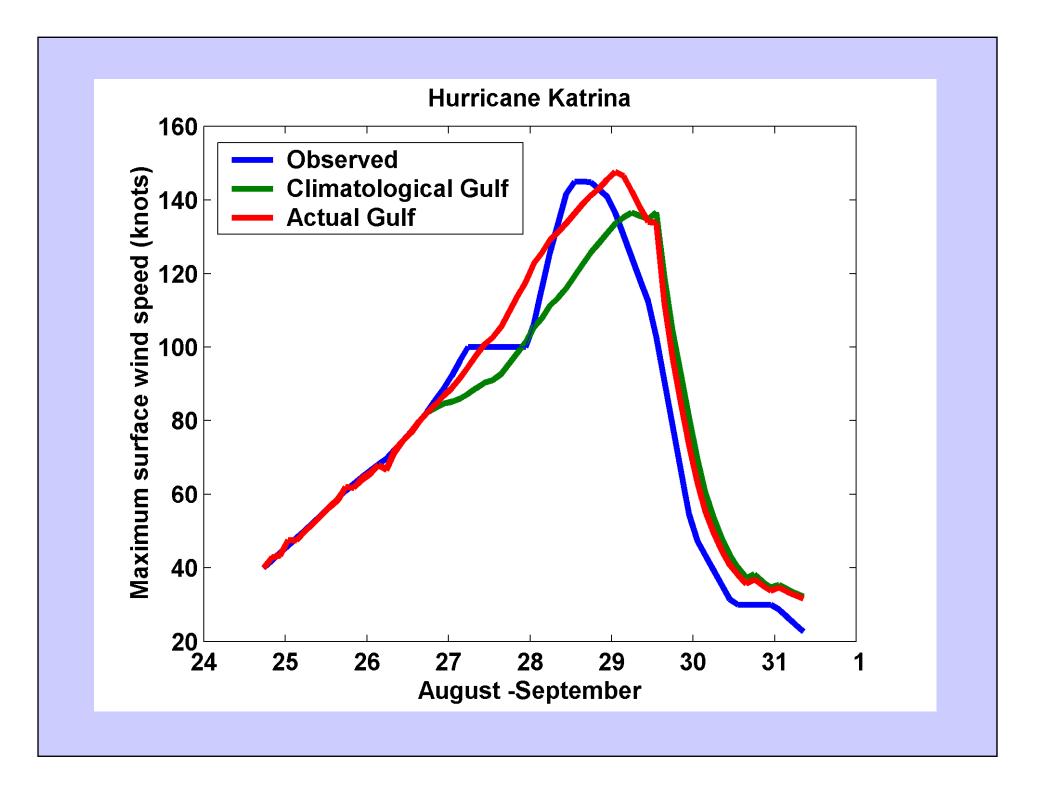
$$PDI \equiv \int_0^\tau V_{max}^3 dt$$

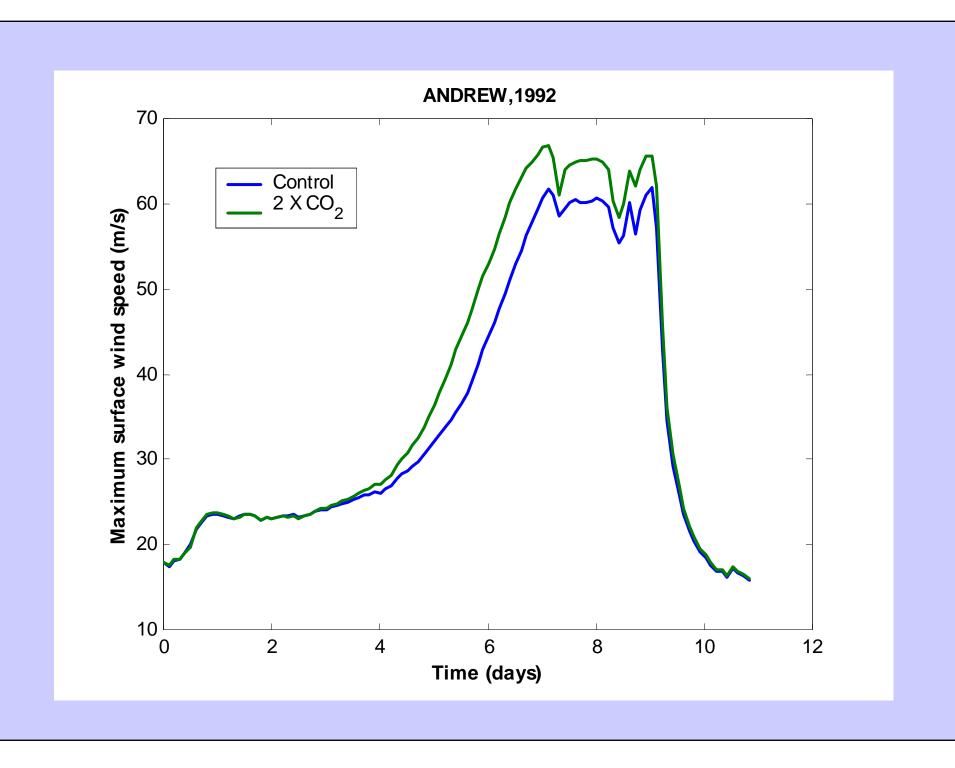
Atlántico



Pacífico NW



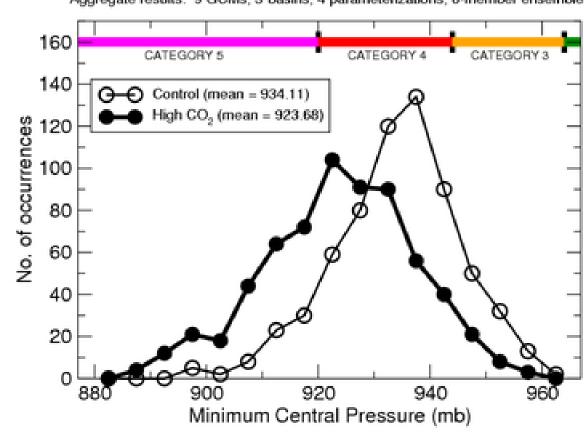


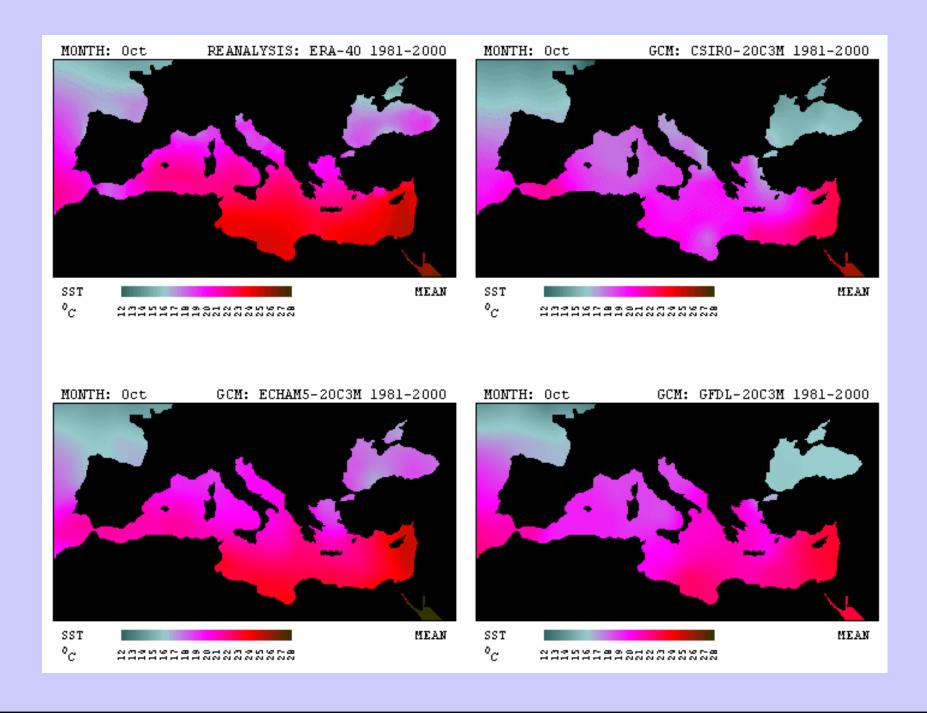


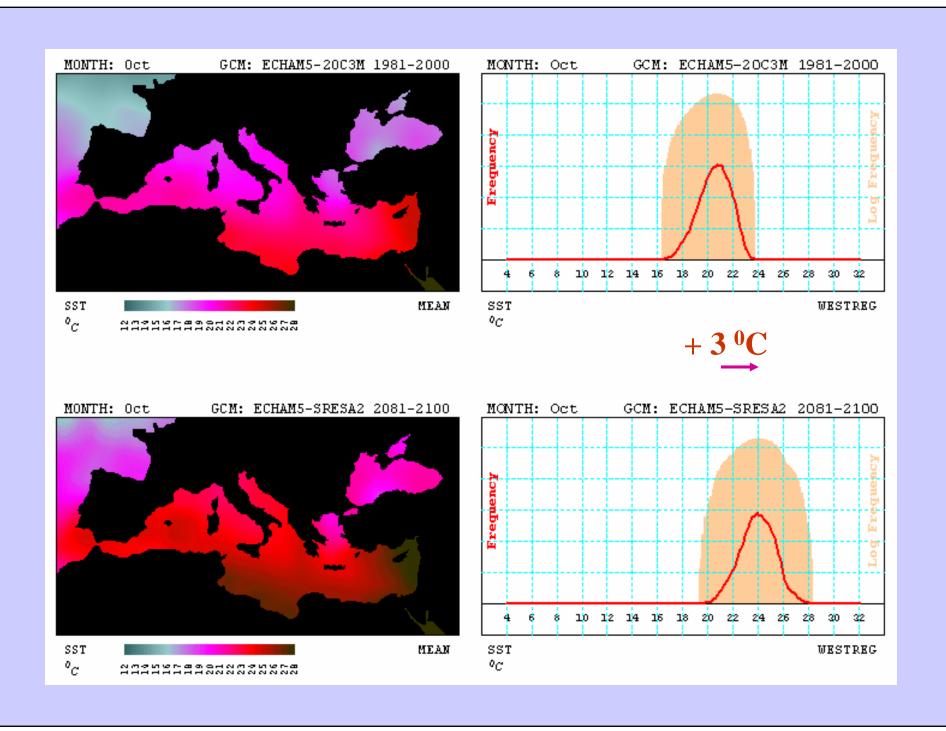
Knutson and Toleya (J. Climate, 2004)



Aggregate results: 9 GCMs, 3 basins, 4 parameterizations, 6-member ensembles







Índice Empírico de Génesis

$$I = \left| 10^{5} \eta \right|^{3/2} \left(\frac{H}{50} \right)^{3} \left(\frac{V_{pot}}{70} \right)^{3} \left(1 + 0.1 V_{shear} \right)^{-2},$$

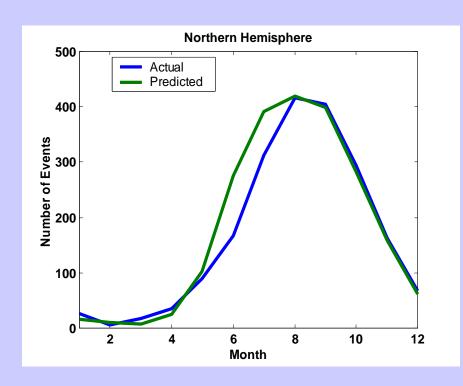
 $\eta \equiv 850 \, hPa \, absolute \, vorticity \, (s^{-1}),$

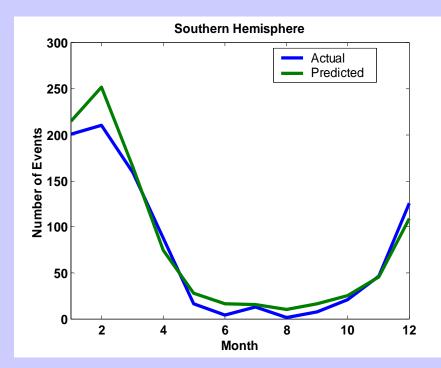
 $V_{pot} = Potential wind speed (ms^{-1}),$

 $H \equiv 600 \, mb \, relative \, humidity (\%),$

$$V_{shear} \equiv \begin{vmatrix} \mathbf{V}_{850} - \mathbf{V}_{250} \end{vmatrix}$$
 (ms⁻¹).

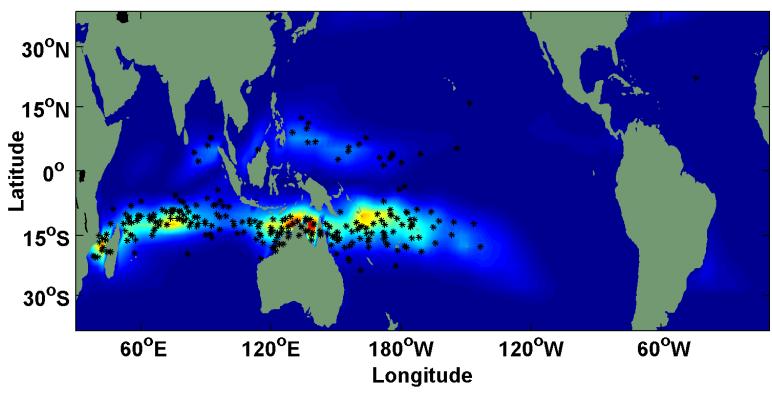
Variabilidad Estacional

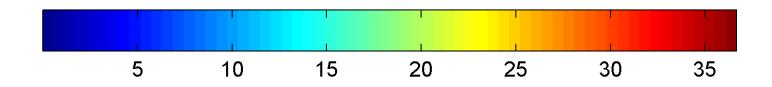




Variabilidad Espacial (H.S)

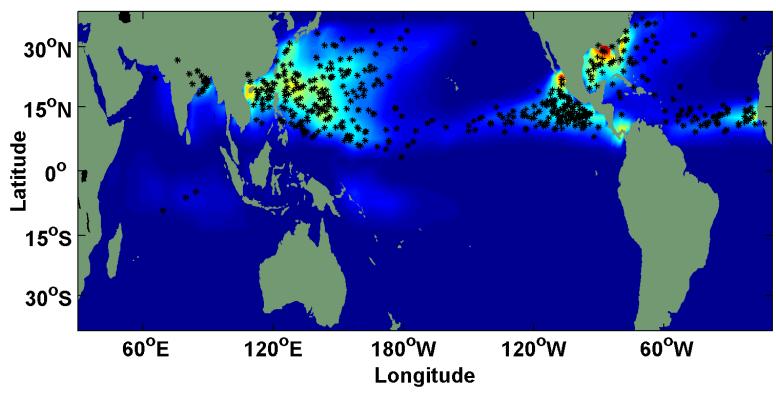


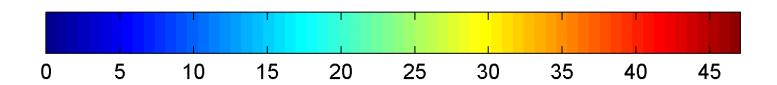


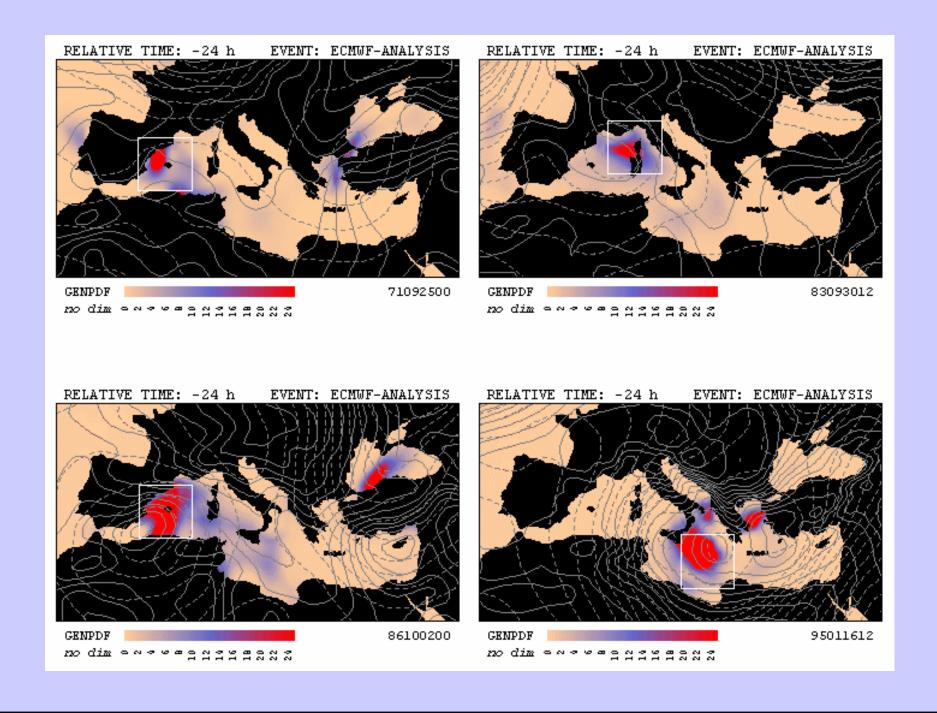


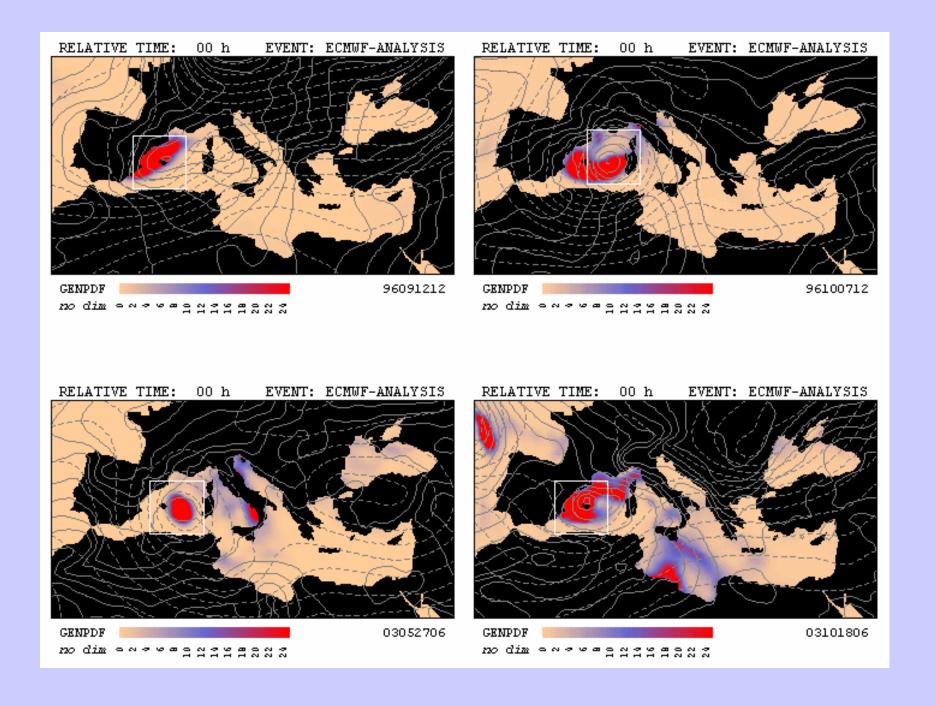
Variabilidad Espacial (H.N)

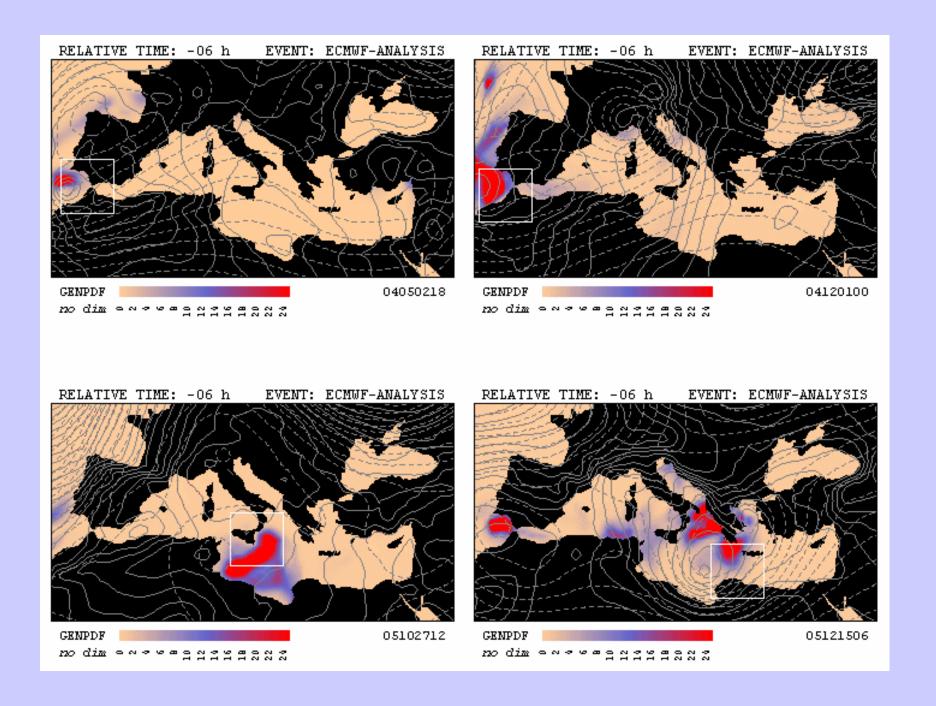
August 1971-2003 (# storms per decade per 2.5° square)

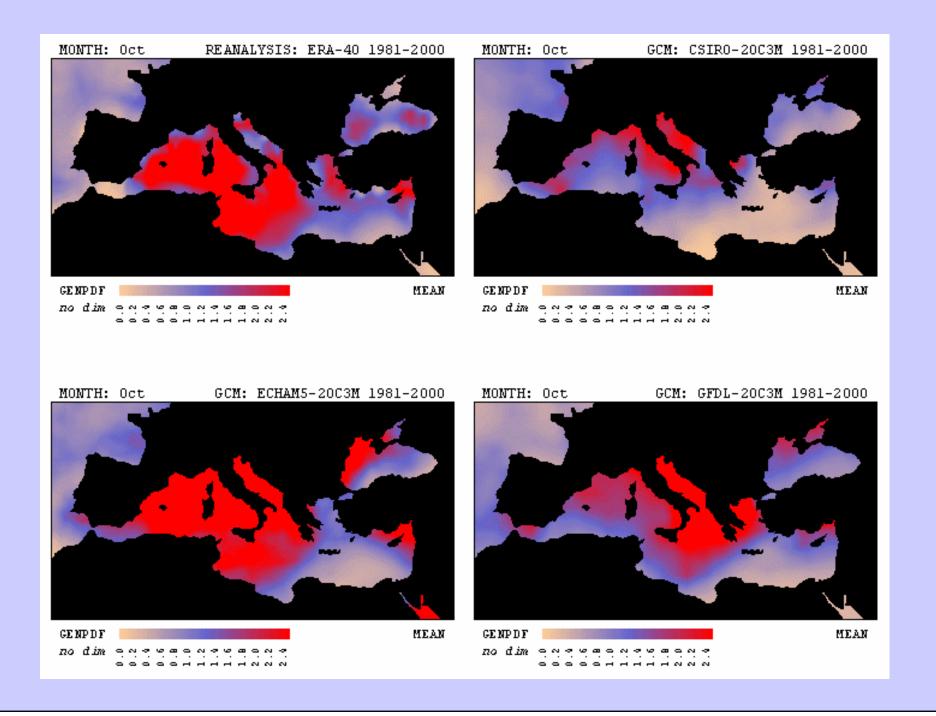


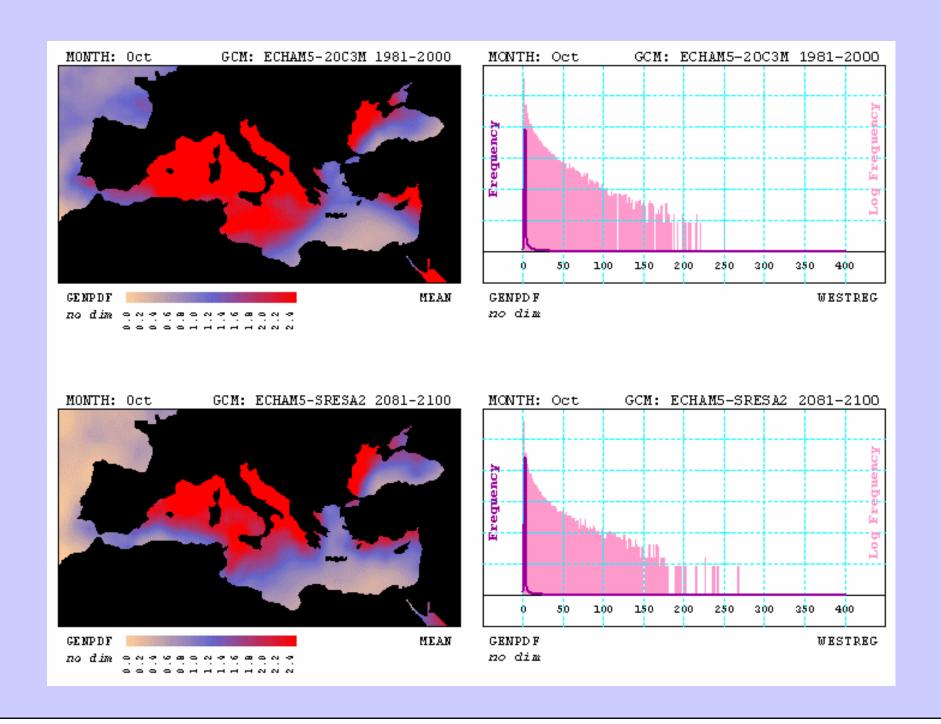




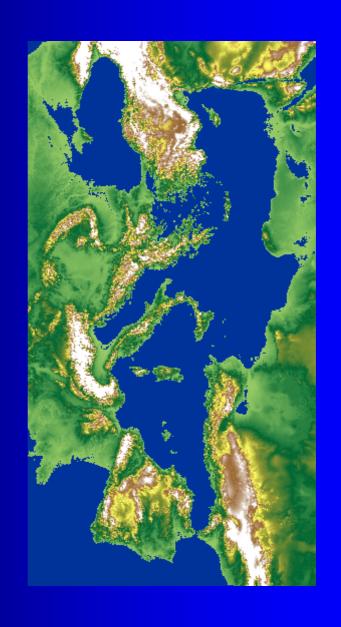








Spouled of the mothods



MOTIVATION

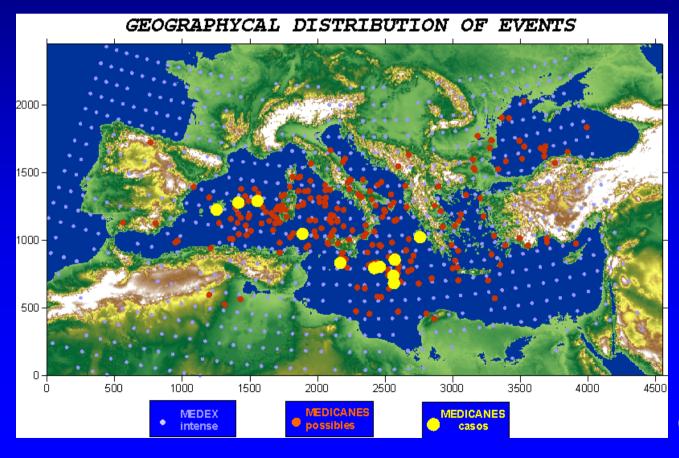
Medicanes are warm-core, surface flux-driven extreme windstorms potentially threatening the islands and coastal areas:

- Are there favoured locations for medicane development ?
- How intense can they become ?
- How could they react in frequency and intensity to global warming ?



MEDICANE RISK???

With an average frequency of only 1-2 events per year and given the lack of systematic, multidecadal databases, an objective evaluation of the long-term risk of medicane-induced winds is impractical with standard methods

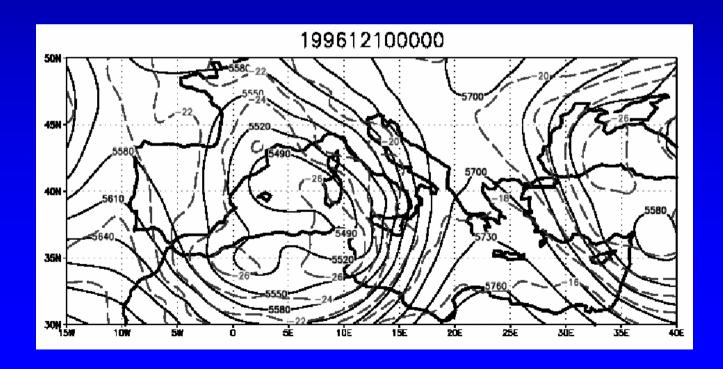


Database from satellite (Tous and Romero, 2012)

APPROACH: Large-scale environmental proxies

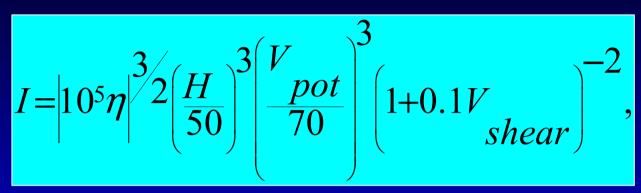
Synoptic analyses of a few studied cases show that an inevitable precursor is the presence of a deep, cut-off, cold-core low in the upper and middle troposphere:

• But the infrequent occurrence of medicanes suggests that additional and very special meteorological conditions are necessary for these storms to occur ...



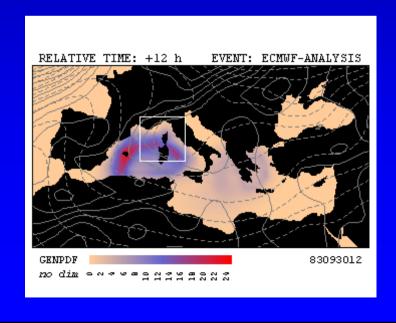
APPROACH: Large-scale environmental proxies

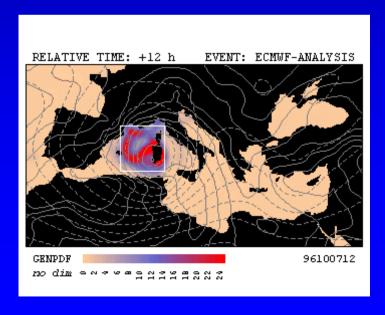
Application of an empirical index of genesis:



GENIX parameter (Emanuel and Nolan, 2004)

 But these environmental proxies behave as necessary but no sufficient ingredients for the successful genesis of a medicane ...

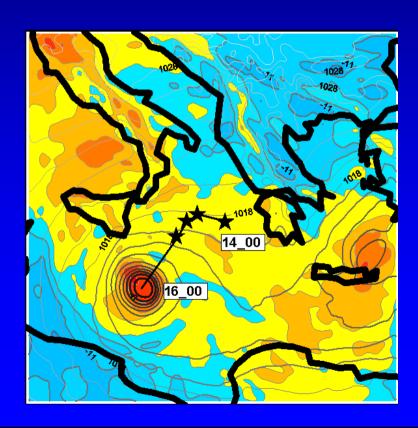


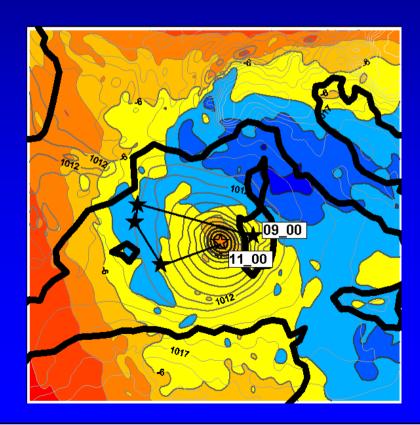


FIRST METHOD: Nested climatic simulations

Detection and tracking of symmetric warm-core cyclonic disturbances generated in mesoscale simulations forced by Reanalysis and GCM data:

• But high computational cost: Limited horizontal resolution; Too few climatic realizations to permit a full sampling of the PDF of storms ...





SECOND METHOD: Statistical-deterministic approach

Developed by Kerry Emanuel and his team in the context of the long-term wind risk associated with tropical cyclones:

- Low-cost generation of thousands of synthetic storms
- Statistically robust assessment of risk (e.g. return periods for winds)
- Genesis: Random draws from observed PDF or Random seeding
- Track: Randomly varying synthetic winds (respecting climatology)
- Environment: Previous winds + monthly-mean thermodynamic fields
- Intensity and radial distribution of winds: CHIPS model

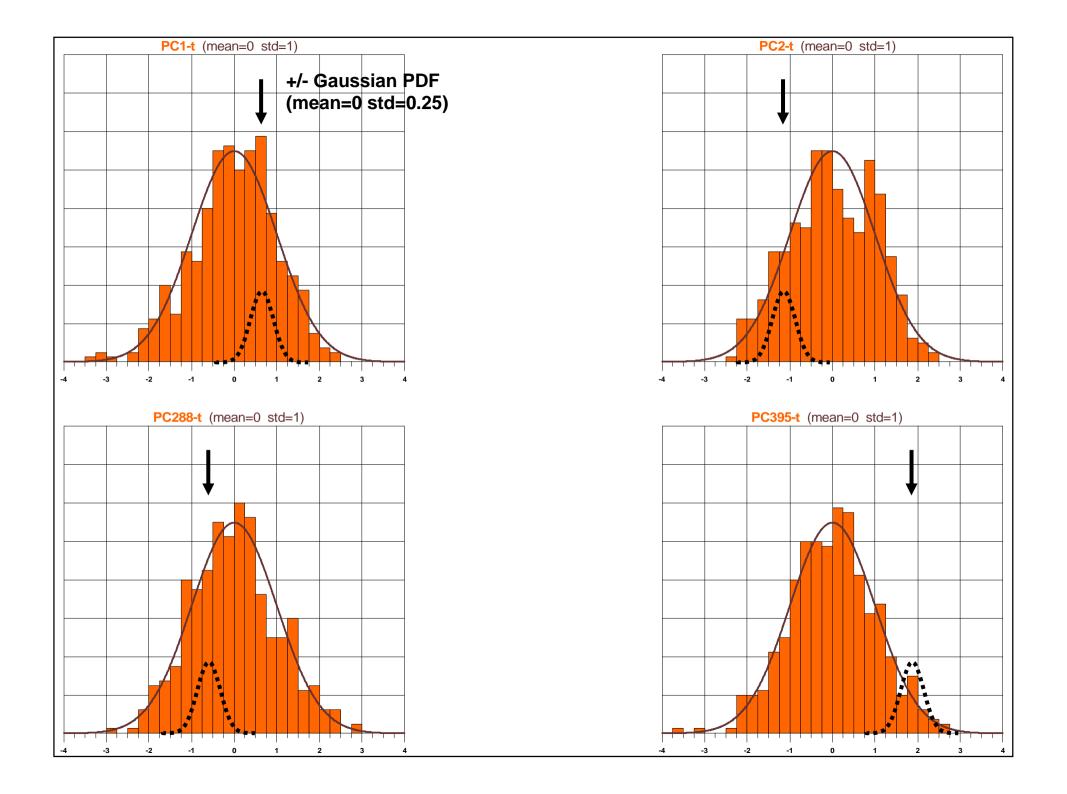


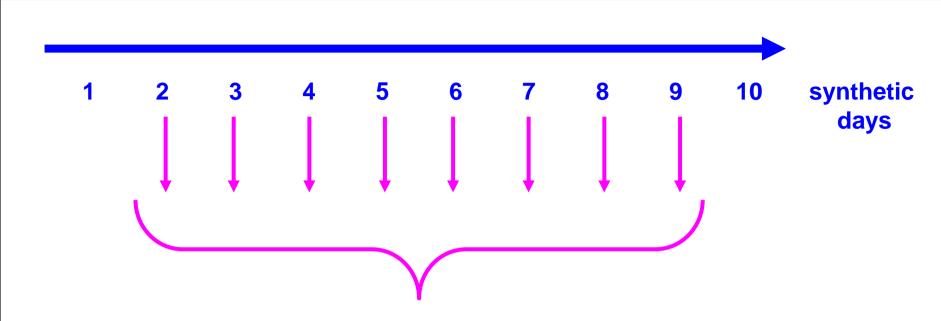
ADAPTATION OF THE SECOND METHOD

The separation of timescales made in the tropics between the synthetic wind field (fast scale) and the thermodynamic environment (slow scale) is not appropriate to represent the movement, growth and decay of mid-latitude weather systems. In addition, the history of medicane genesis is far too sparse to form a reasonable PDF of genesis, and random seeding would be very inefficient:

- For each month, decomposition through PCA of 10-day synoptic evolutions of z250, z850, T600, R600 and PINT into the new space of independent PCs
- Random selection + random perturbation of the set of PCs
- This perturbed set of PCs is converted back into physical space
- This is tantamount to generating 10-day sequences of spatiotemporal coherent z250, z850, T600, R600 and PINT synthetic fields which also respect their mutual covariances

Potential Genesis: Based on the GENIX parameter

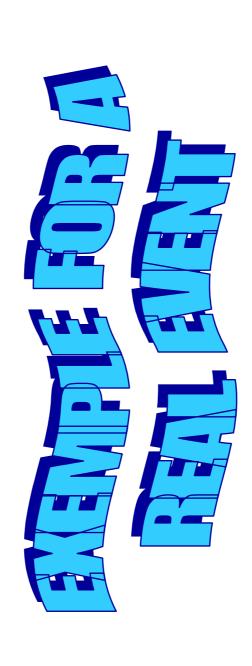


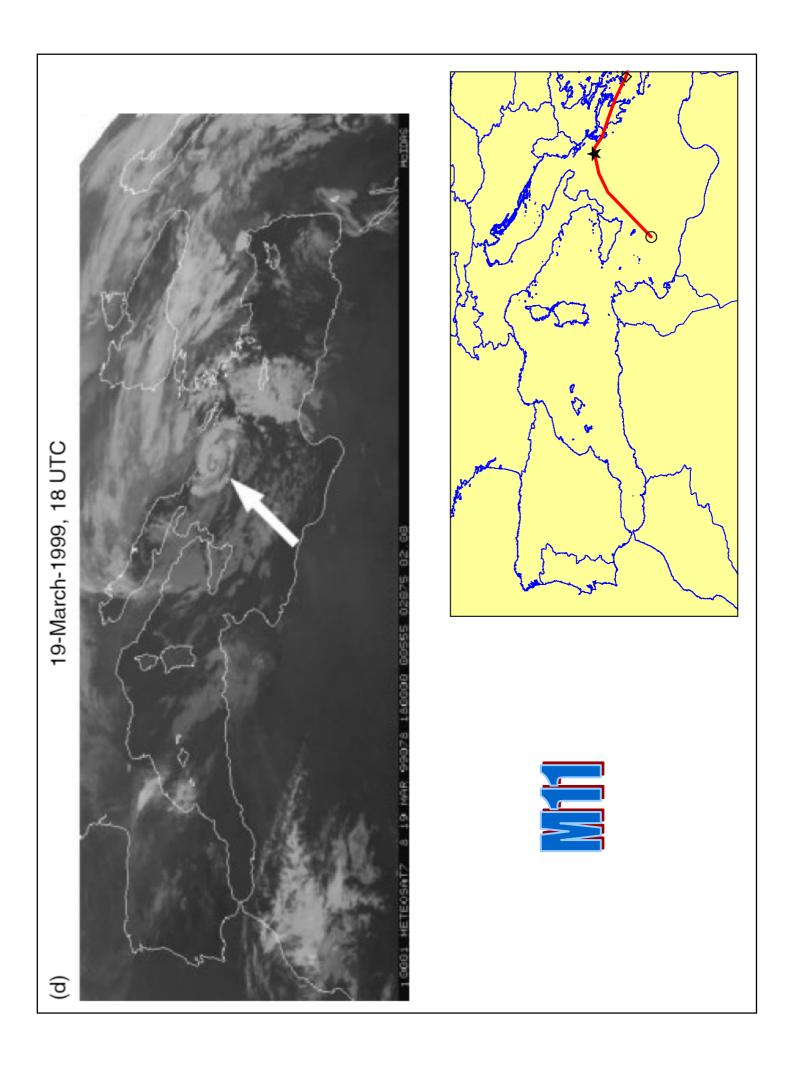


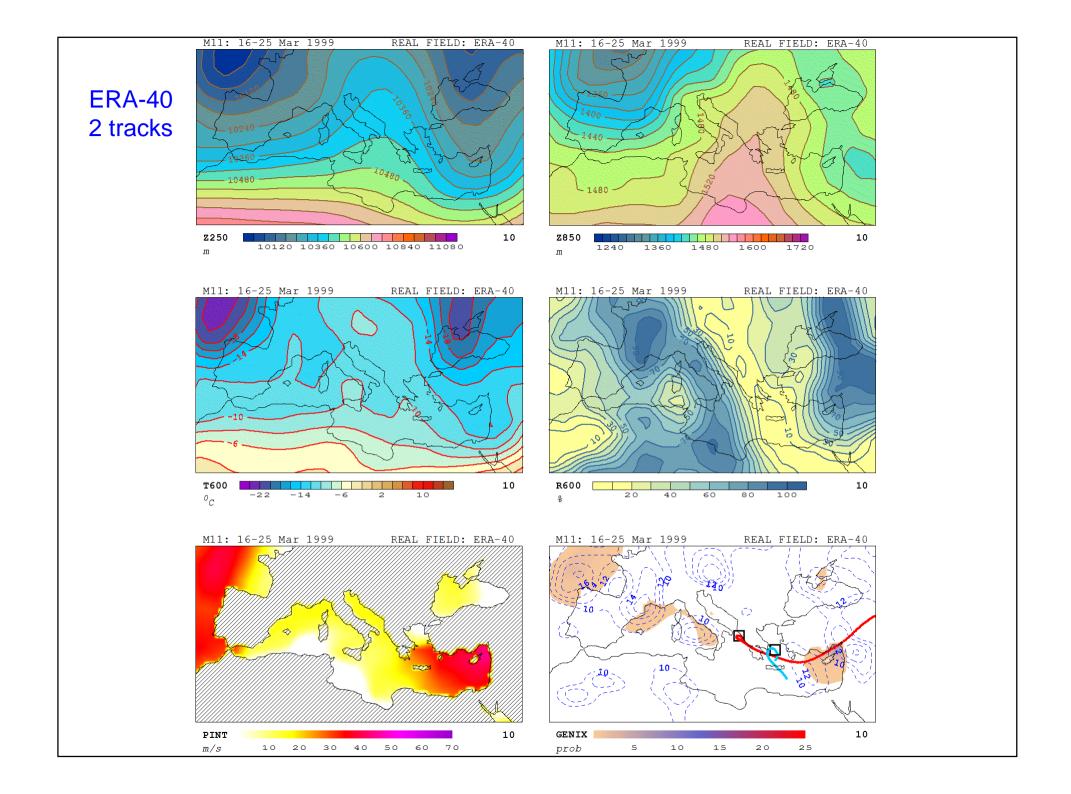
OPEN-SEA POINT + MAX OF GENIX > 20 + ABS VOR > 10 units ???

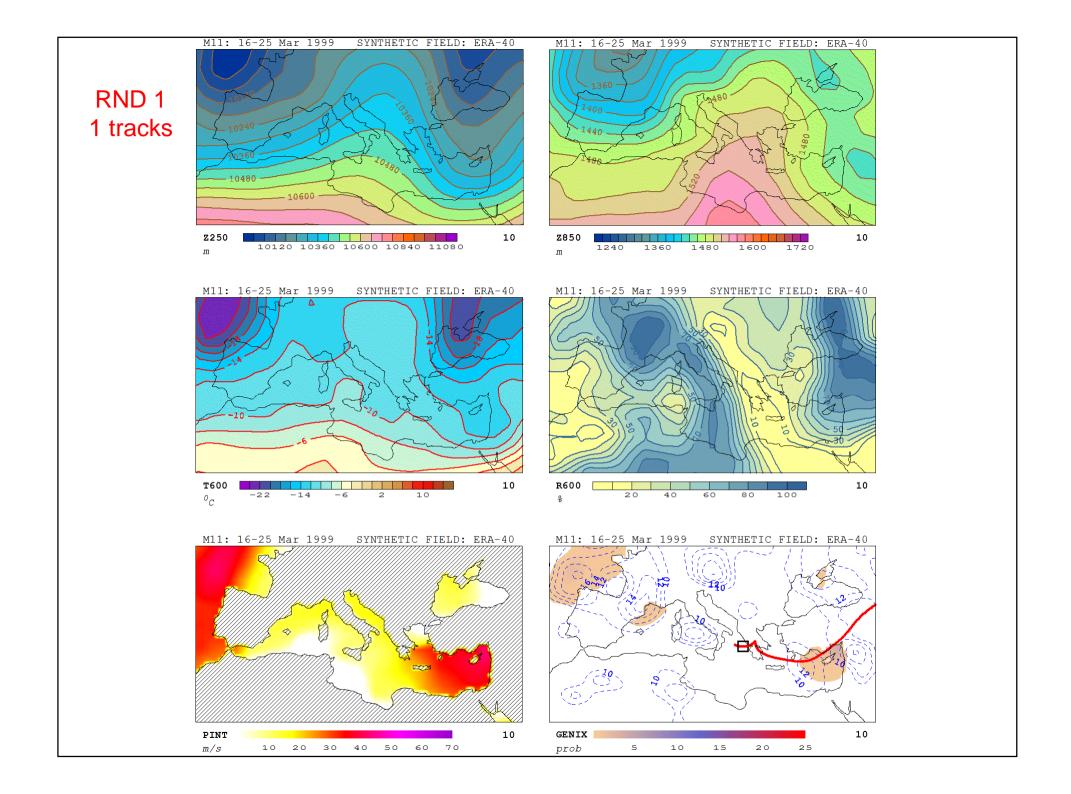


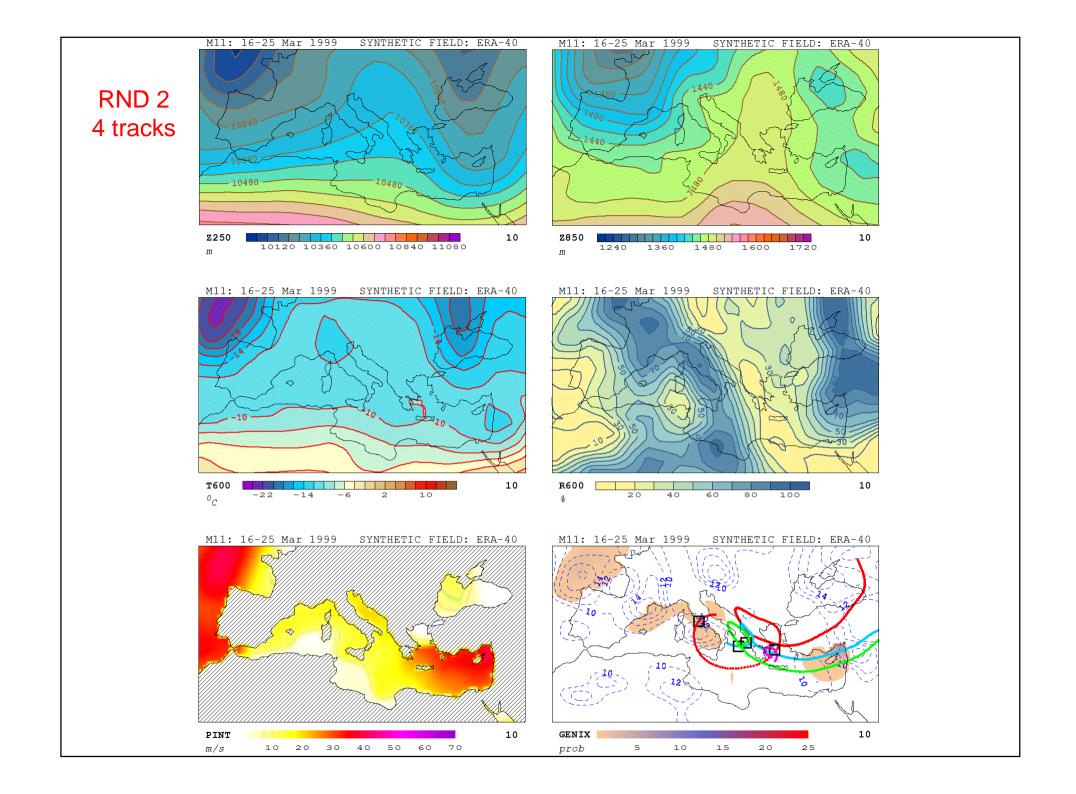
$$\begin{cases} u_{track} = \alpha \ u_{850} + (1 - \alpha) \ u_{250} \\ v_{track} = \alpha \ v_{850} + (1 - \alpha) \ v_{250} + v_{\beta} \end{cases} \qquad \alpha = 0.8 \quad v_{\beta} = 2.5 \text{ m/s}$$

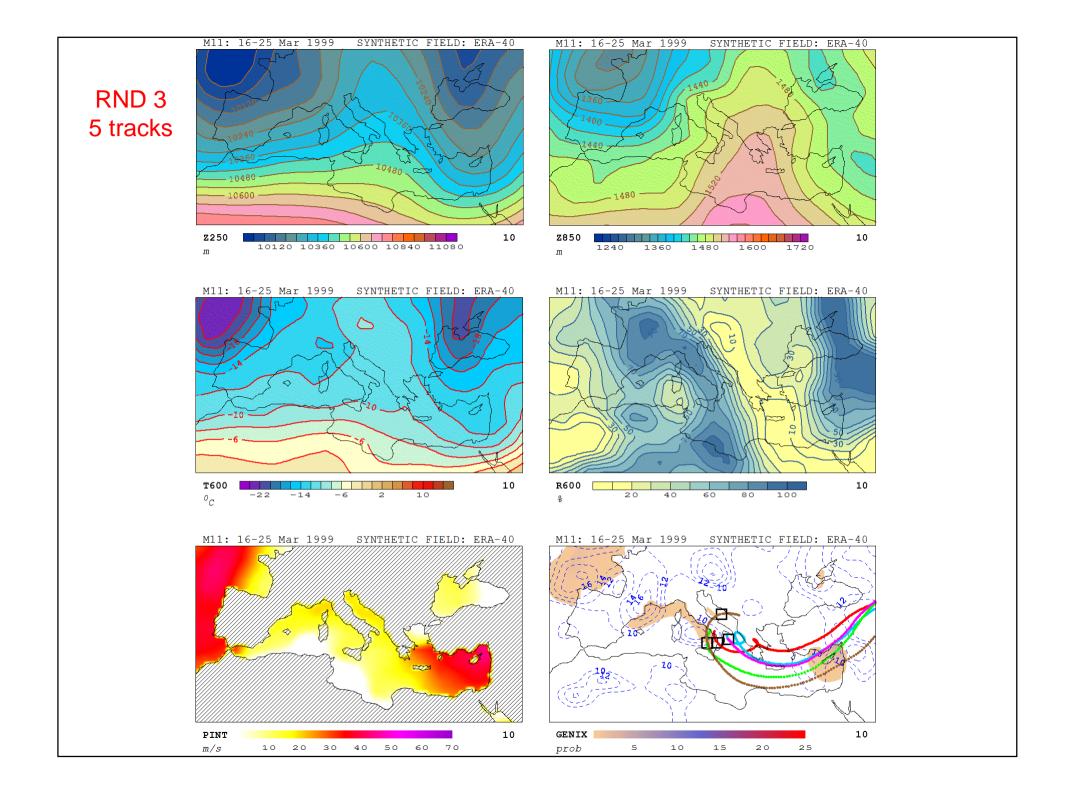


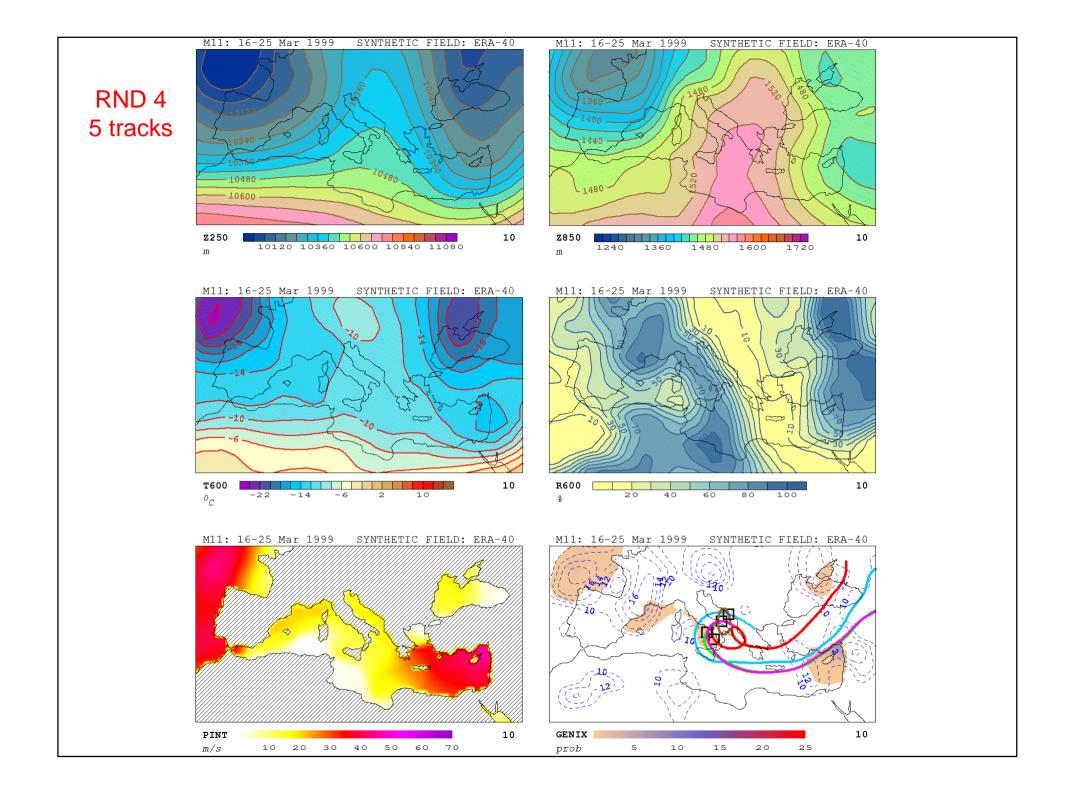


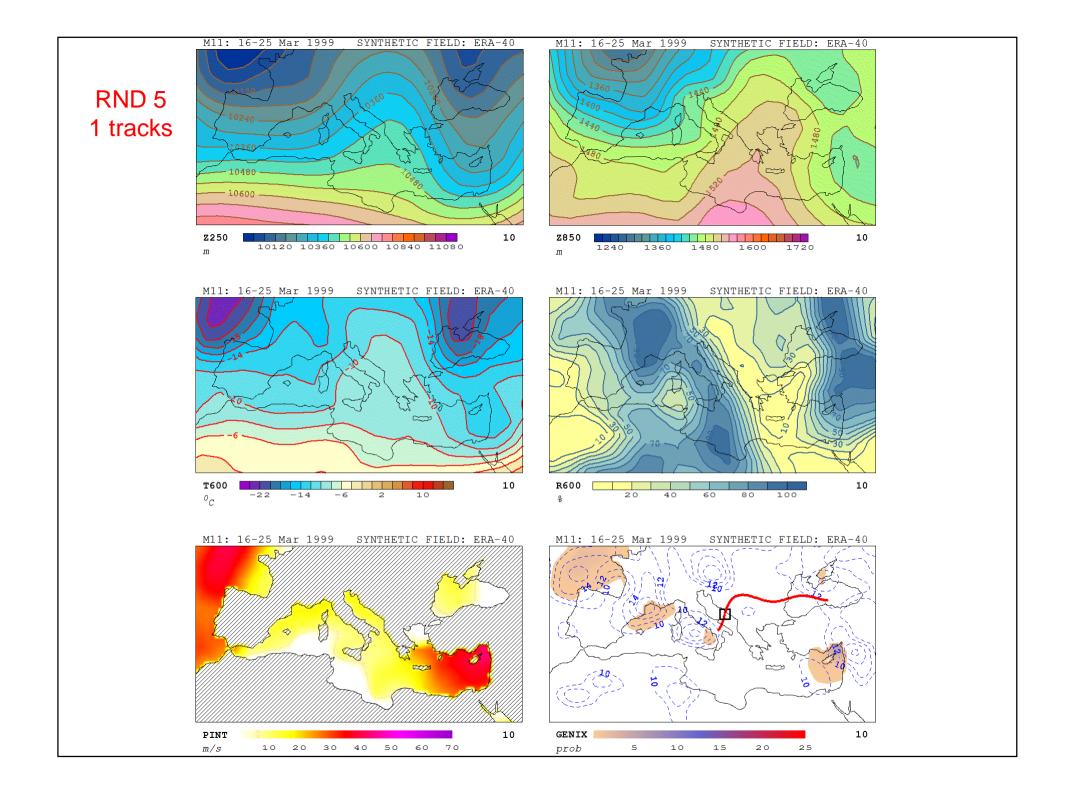








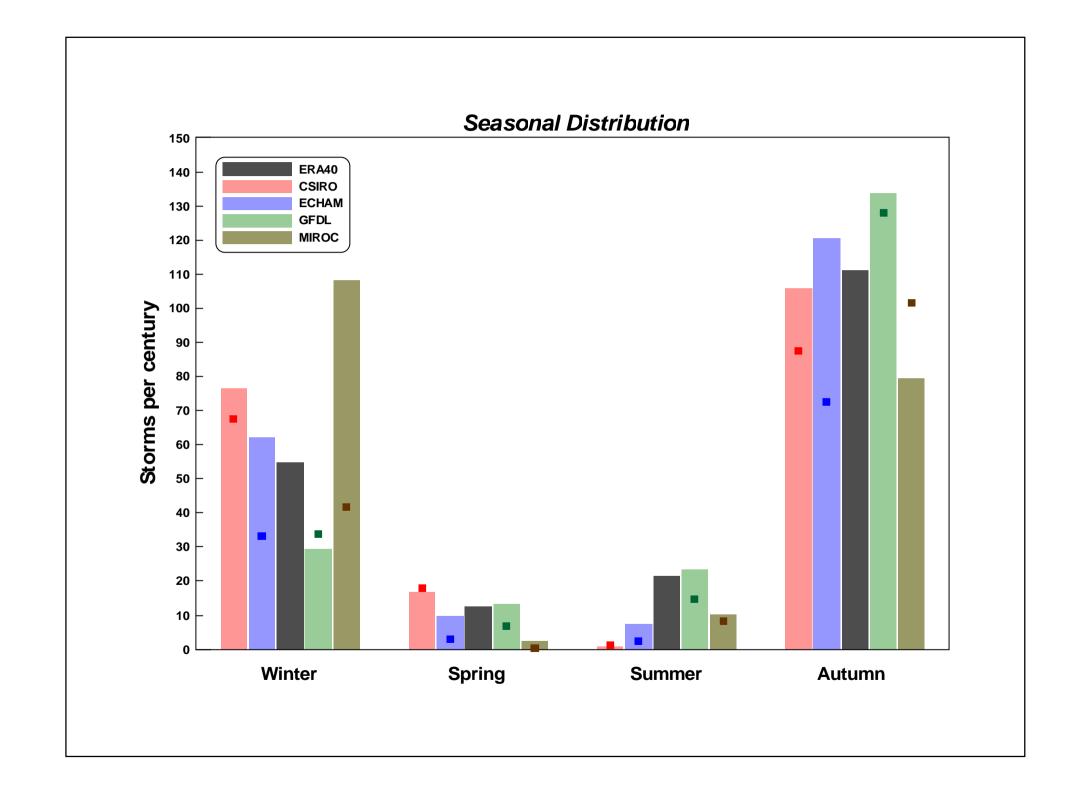


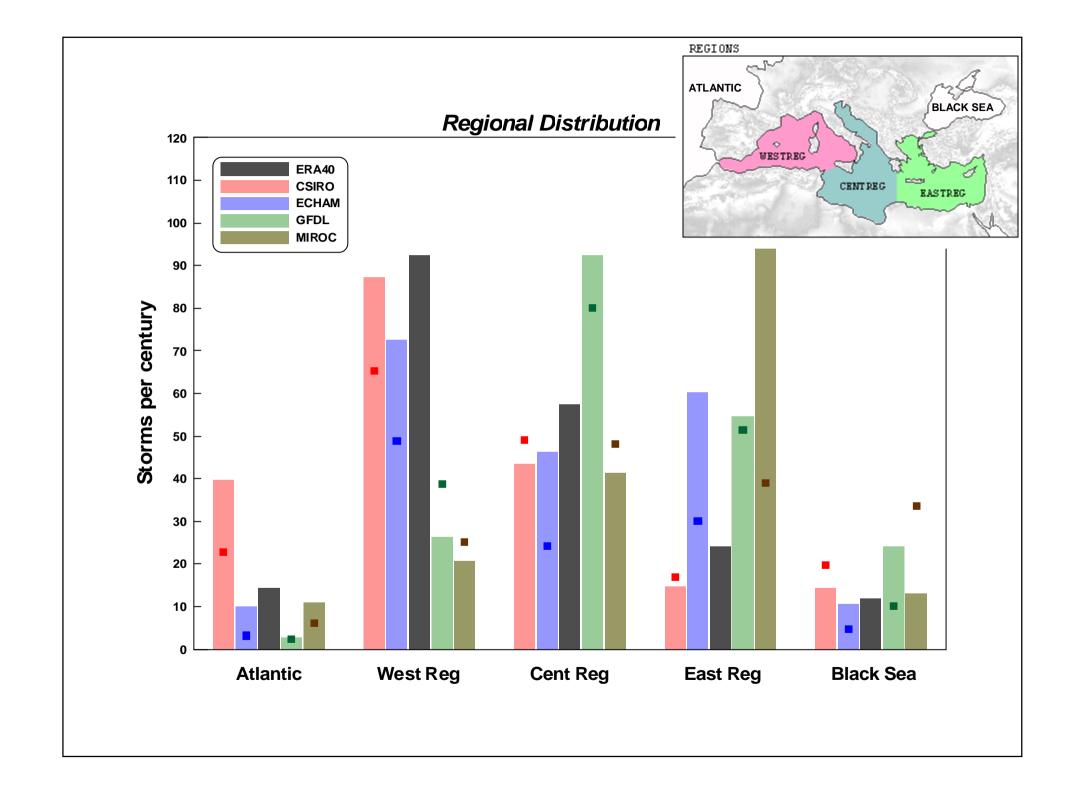


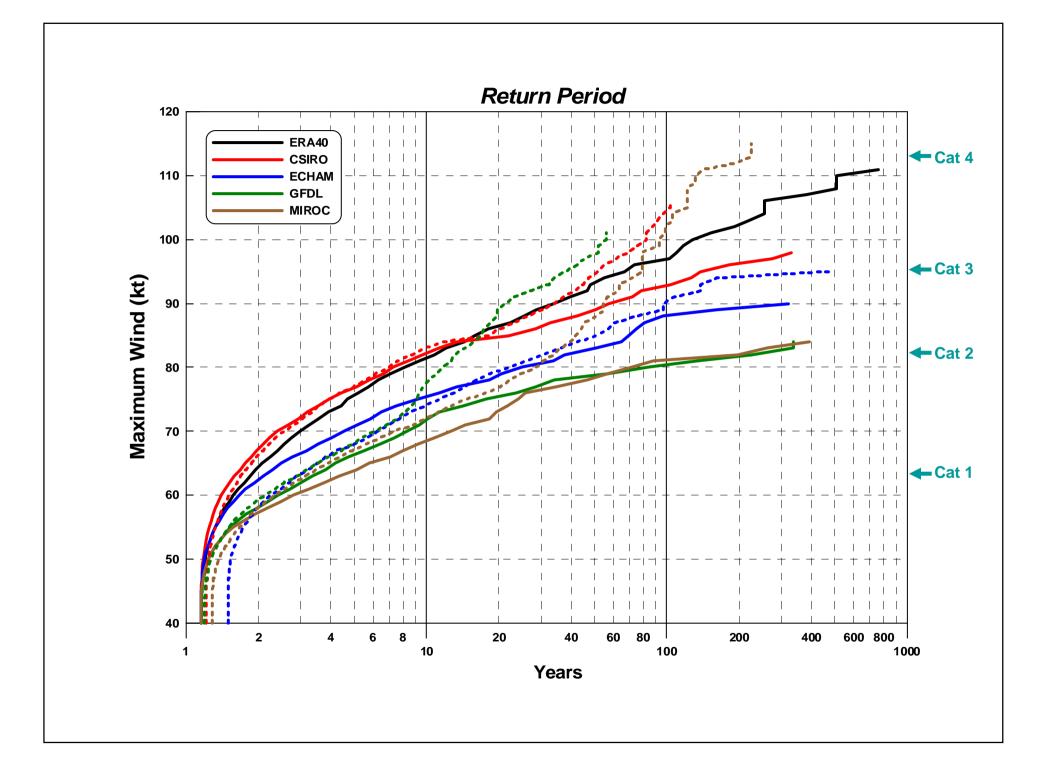
APPLICATION OF THE SECOND METHOD

We synthetically generate a total of ~15000 potential tracks for each climate/model. These are simulated with CHIPS and checked for intensification above TS category (34 kt):

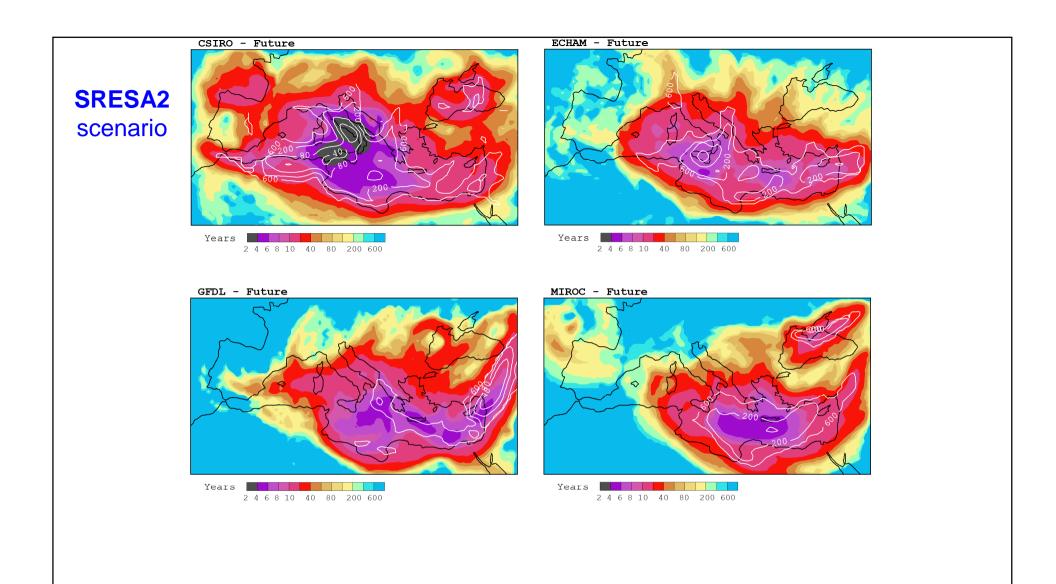
Climate Scenario	Reanalysis or GCM	Successful Storms	Storms per century
	ERA40	3048	200
	CSIRO	3286	200
PRESENT	ECHAM	1924	200
1981 – 2000	GFDL	1343	200
	MIROC	1567	200
	CSIRO	2857	174
FUTURE	ECHAM	1072	111
2081 – 2100	GFDL	1226	183
SRES A2	MIROC	2389	152



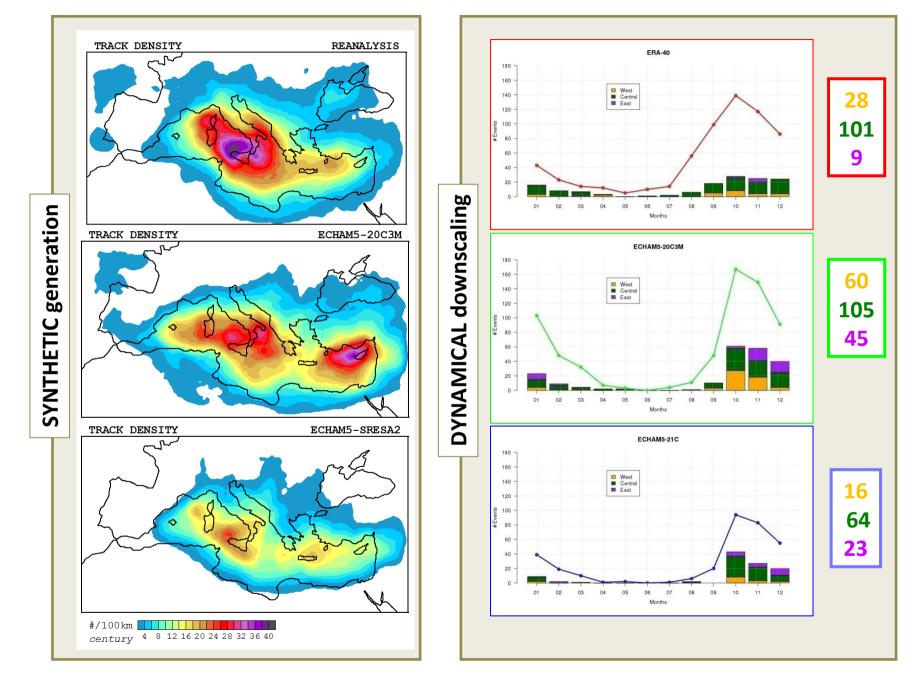




CSIRO - Present ECHAM - Present **20C3M** scenario Years 2 4 6 8 10 40 80 200 600 2 4 6 8 10 40 80 200 600 GFDL - Present MIROC - Present Years 2 4 6 8 10 40 80 200 600 Years 2 4 6 8 10 40 80 200 600 ERA40 - Present 2 4 6 8 10 40 80 200 600



COMPARISON OF BOTH METHODS



CONCLUSIONS

- The statistical-deterministic approach is a good alternative to computationally expensive classical methods (e.g. dynamical downscaling of medicanes), with the extra benefit of producing statistically large populations of events
- We attained unprecedented medicane-wind risk maps for the Mediterranean region
- General agreement with the "known" phenomenology of medicanes in the current climate (e.g. maximum in the cold season and central Mediterranean) and between both methods
- In spite of some geographical uncertainties, GCMs tend to project fewer medicanes at the end of the century compared to present but a higher number of violent storms, suggesting an increased probability of major economic and social impacts as the century progresses

